

Solutions to Laplace's Eqⁿ

- We will look @ cartesian + spherical coordinates

- Approach: sepⁿ of variables

- Point \ni MATH Green's 3rd identity \implies analytical continuation
 IF $u, \partial u / \partial n$ known over surface bounding
 a source-free region, R , then u can be
 calculated anywhere in the region.

but HOW? ie there is a unique solⁿ
for u .

$\times \times \times \times \times$ observations

calculate u here

OR HERE



$$\Delta p = p' - p = 0 \text{ elsewhere}$$

$$\nabla^2 u = 0 \implies \frac{\partial^2 u}{\partial x^2} + \frac{\partial^2 u}{\partial y^2} + \frac{\partial^2 u}{\partial z^2} = 0$$

choose a solⁿ of the form

$$u(u, v, x, y, z) = X(u, x) Y(v, y) Z(u, v, z)$$

\downarrow this is only 1 possible solⁿ & is
 specified by choice of u, v

\downarrow separation constants

Calculate $\frac{1}{u} \nabla^2 u$ ($= 0$)

Substitute

$u = xyz$ &

divide by xyz

$$= \frac{1}{xyz} \frac{\partial^2 (xyz)}{\partial x^2} + \dots$$

$$= \frac{1}{x} \frac{\partial^2 y}{\partial x^2} + \frac{1}{y} \frac{\partial^2 x}{\partial y^2} + \frac{1}{z} \frac{\partial^2 z}{\partial z^2}$$

$$\Rightarrow \frac{1}{x} \frac{\partial^2 x}{\partial x^2} = -\frac{1}{y} \frac{\partial^2 y}{\partial y^2} - \frac{1}{z} \frac{\partial^2 z}{\partial z^2}$$

\downarrow \downarrow \downarrow
 $f(x)$ $f(y)$ $f(z)$

\Rightarrow must be a constant
 set the constant = $-u^2$

LHS \Rightarrow $\boxed{\frac{\partial^2 x}{\partial x^2} + u^2 x = 0}$ --- ①

similarly
 by substituting

$$\frac{1}{y} \frac{\partial^2 y}{\partial y^2} = u^2 - \frac{1}{z} \frac{\partial^2 z}{\partial z^2} = \text{const.} = -v^2$$

\downarrow \downarrow
 $f(y)$ $f(z)$

$$\Rightarrow \boxed{\frac{\partial^2 y}{\partial y^2} + v^2 y = 0}$$
 --- ②

and now

$$\frac{1}{z} \frac{\partial^2 z}{\partial z^2} = u^2 + v^2$$

$$\Rightarrow \boxed{\frac{\partial^2 z}{\partial z^2} - (u^2 + v^2) z = 0}$$
 --- ③

Equations ① - ③ have solutions

$$\begin{aligned} X(u, x) &= a_1(u) e^{iux} + b_1(u) e^{-iux} & u \geq 0 \\ &= c_1(u) e^{iux} & -\infty < u < \infty \\ Y(v, y) &= c_2(v) e^{ivy} & -\infty < v < \infty \\ Z(u, v, z) &= c_3(u, v) e^{\pm(u^2+v^2)^{1/2} z} \end{aligned}$$

so a solution to Laplace's eqⁿ is

$$u = X(u, x) Y(v, y) Z(u, v, z)$$

$$\textcircled{4} \quad u(x, y, z) = \left[A(u, v) e^{-(u^2+v^2)^{1/2} z} + B(u, v) e^{(u^2+v^2)^{1/2} z} \right] \underbrace{e^{iux} e^{ivy}}_{e^{i(ux+vy)}}$$

THIS IS ONE SOLUTION, specified by the values of u and v .

Obtain the general solution by adding the contributions for all possible values of u and v .

TWO CASES

A// u, v take discrete values

$$\left. \begin{aligned} u &= 2\pi n / L \\ v &= 2\pi m / L' \end{aligned} \right\} \begin{array}{l} n, m \text{ are integers} \\ L, L' \text{ period} \end{array}$$

corresponds to periodic solⁿ & the addⁿ is a summation

$$\sum_m \sum_n$$

B// u, v vary continuously

Contribution to the general solⁿ from param values btwn

u and $u + \delta u$, and

v and $v + \delta v$

is expressed by weighting fⁿs $F(u, v)$, $G(u, v)$

$$\begin{aligned} \text{where} \quad \frac{1}{2\pi} \int F(u, v) \delta u \delta v &= A(u, v) & 2\pi \text{ convenient} \\ \frac{1}{2\pi} \int G(u, v) \delta u \delta v &= B(u, v) & \rightarrow \end{aligned}$$

addition of all possible solids is now an integral over u & v rather than a discrete summation

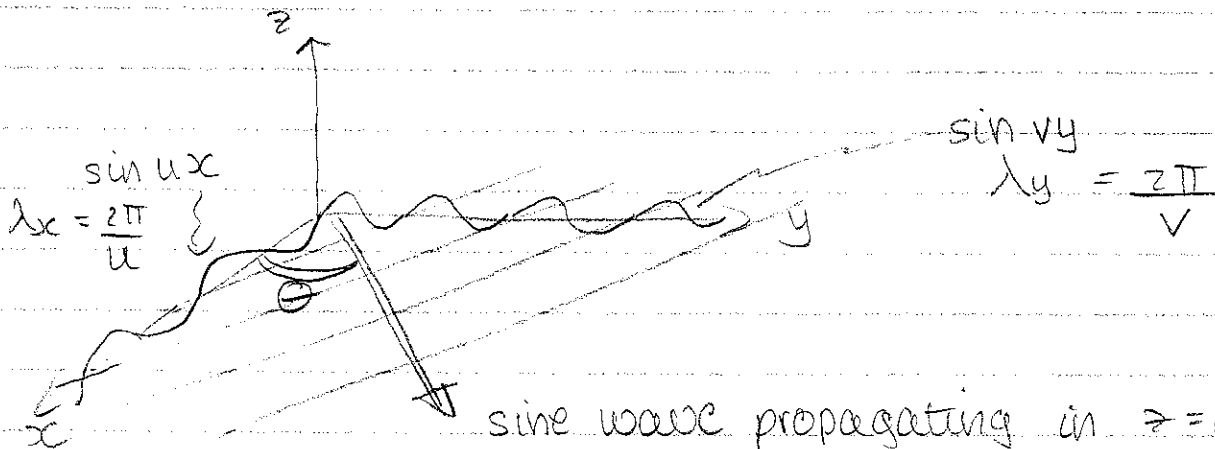
$$u(x, y, z) = \frac{1}{2\pi} \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} e^{i(ux+vy)} \left\{ F(u, v) e^{-(u^2+v^2)^{1/2} z} + G(u, v) e^{(u^2+v^2)^{1/2} z} \right\} du dv$$

PHYSICAL INTERPRETATION

when $z=0$

$$u(x, y, z) = \frac{1}{2\pi} \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} (F(u, v) + G(u, v)) e^{i(ux+vy)} du dv$$

But we have seen this before: $z=0$ FT!



sine wave propagating in $z=0$ plane
direction defined by u, v $\theta=0 \Rightarrow$

$$\tan \theta = \frac{v}{u}$$

check $v=0 \Rightarrow$ only in \hat{x} dirⁿ

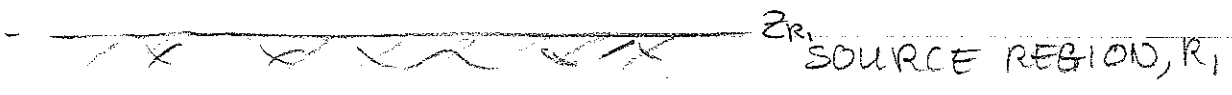
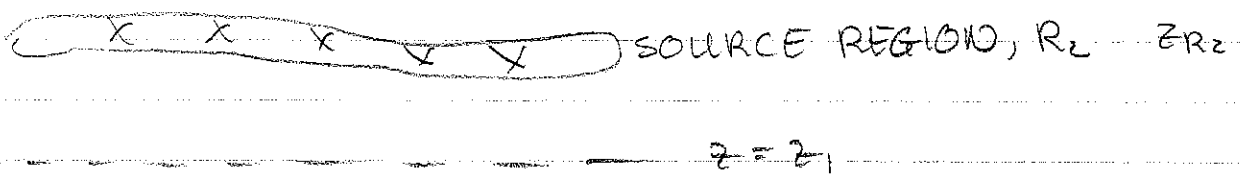
$u=0 \Rightarrow \theta=90^\circ \Rightarrow$ only in \hat{y}

$F(u, v)$ and $G(u, v)$ are the amplitude of the sinusoid on $z=0$

Consider $z \neq 0$

$F(u, v)$ term	\downarrow exponentially	} for $z > 0$
$G(u, v)$ term	\uparrow exponentially	

Consider obs'n's on a plane $z = z_1$



$F(u, v) e^{-(u^2+v^2)^{1/2}(z_1 - z_{R1})}$ amplitude of sinusoidal wave's ($e^{i(u x + v y)}$)
 whose sources lie below z_1 , (@ ~~z_{R1}~~ z_{R1})
 $G(u, v) e^{(u^2+v^2)^{1/2}(z_1 - z_{R2})}$ - - - - - sources above z_1 , (@ z_{R2} here)

This is the ESSENCE OF UPWARD / DOWNWARD CONTINUATION

Partie in sph.
geom.

$$\textcircled{5} \quad \begin{matrix} \varphi_B \\ \vec{g} \\ N \end{matrix} \quad (g_x, g_y, g_z)$$

$$\begin{matrix} \varphi_B \\ \vec{B} \\ \vec{m} \end{matrix}$$

ρ, H, v_s, v_p

why?

because integrand's orthog
so it's an expansion
in terms of basic fns

$$|k| = \sqrt{k_x^2 + k_y^2}$$

$$= \frac{1}{2\pi} \sqrt{u^2 + v^2}$$

$$\lambda = \frac{1}{|k|} = \frac{2\pi}{\sqrt{u^2 + v^2}} = \frac{2\pi}{2\pi \sqrt{k_x^2 + k_y^2}}$$

$$= \frac{1}{\sqrt{k_x^2 + k_y^2}}$$

FROM LAST TIME

$\nabla^2 U = 0$

$$\Rightarrow U(x, y, z) = \frac{1}{2\pi} \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} e^{i(ux+vy)} \left\{ F(u, v) e^{-(u^2+v^2)^{1/2} z} + G(u, v) e^{(u^2+v^2)^{1/2} z} \right\} du dv$$

potential in (x, y, z) system

u, v spatial frequency

$u = 2\pi k_x \quad k_x = \text{wave \# in } x \text{ dir}^m = \frac{1}{\lambda_x} \rightarrow \text{wavelength}$
 $v = 2\pi k_y \quad k_y = 1/\lambda_y$

wave w/ $\lambda = \frac{1}{|k|}$

$\Theta = \tan^{-1}(v/u)$

Important stuff to remember

- ① spatial derivatives (and any linear combo of them) of u also satisfy Laplace's eqⁿ + can be represented by ① or for later lateral field anomalies.
- ② If sources for $u, \vec{g}, \vec{B} \dots$ are below $z=0$ then $G(u, v) = 0$ (if not $u \rightarrow \infty$ as $z \rightarrow \infty$)
Conversely $F(u, v) = 0$ for sources above $z=0$.
- ③ same ① on the plane $z=0$ (or on any const z plane) looks like a 2-D FFT
- ④ surface harmonic (varⁿ of u on const z) ; solid harmonic \updownarrow

eq on $z=0$

$$U(x, y, 0) = \frac{1}{2\pi} \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} e^{i(ux+vy)} F(u, v) du dv$$

sources $z < 0$

so, now obvious how to calculate $F(u, v)$!

Take $\mathcal{F}[u(x, y, 0)] \rightarrow F(u, v) = \frac{1}{2\pi} \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} u(x, y, 0) e^{-i(ux+vy)} dx dy$

⑤ see ←

could have used z_0 — — — — — \rightarrow extra term in here $e^{-\sqrt{u^2+v^2} z_0}$

Can calculate u at any z

an alternative approach (why FTs are so slick...)

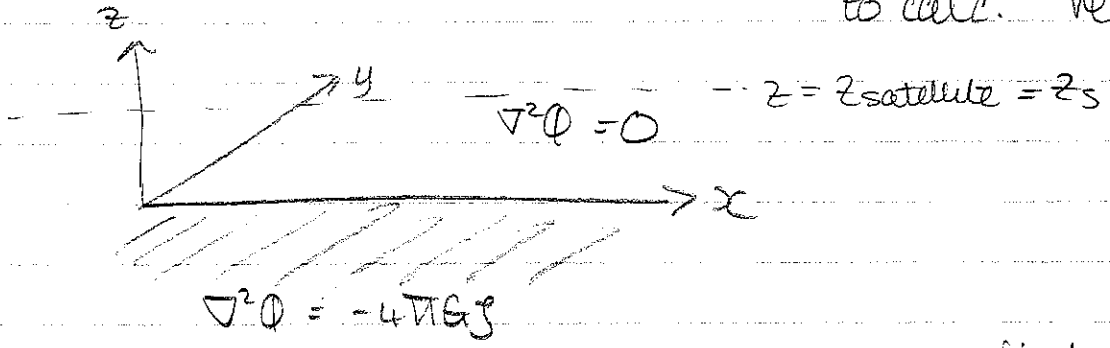
consider satellite measuring Φ (the 'disturbing' or 'anomalous' potential)

$$\Phi = U - U_0$$

total pot^l
reference pot^l

or something that gives us Φ

going to talk ~ how to calc. 'ref' pot's



sources below $\rho = 0$

Assume we can use flat earth approxⁿ

$$\nabla^2 \Phi = 0 \Rightarrow \frac{\partial^2 \Phi}{\partial x^2} + \frac{\partial^2 \Phi}{\partial y^2} + \frac{\partial^2 \Phi}{\partial z^2} = 0 \quad \text{--- (A)}$$

Need 6 B.C.s to solve (A)

$$\left. \begin{aligned} \lim_{|x| \rightarrow \infty} \Phi &= 0 & \lim_{|y| \rightarrow \infty} \Phi &= 0 & \lim_{z \rightarrow \infty} \Phi &= 0 \end{aligned} \right\} \underline{5 \text{ B.C.s}}$$

The 6th one comes from $\Phi(x, y, z_0)$ or $\partial\Phi/\partial z (= -\Delta g)$ on z_0

Before we used sepⁿ of variables

BUT

we can use FTs (to get an algebraic eqⁿ in k_x, k_y)

$$F(\vec{k}) = \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} f(\vec{x}) e^{-2\pi i (\vec{k} \cdot \vec{x})} d^2 \vec{x}$$

$$f(\vec{x}) = \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} F(\vec{k}) e^{2\pi i (\vec{k} \cdot \vec{x})} d^2 \vec{k}$$

where $\vec{x} = (x, y)$ is posⁿ vector

$\vec{k} = (k_x, k_y)$ is wave number vector

and $\vec{k} \cdot \vec{x} = k_x x + k_y y$

Now F.T the P.E.

Recall $\mathcal{F} \left[\frac{\partial \phi}{\partial x} \right] = -i 2\pi k_x \mathcal{F}[\phi]$
 $\mathcal{F} \left[\frac{\partial^2 \phi}{\partial x^2} \right] = -4\pi^2 k_x^2 \mathcal{F}[\phi]$

and likewise for y

Write $\mathcal{F}[\phi(x, y, z)] = \phi(\vec{k}, z)$

So $\frac{\partial^2 \phi}{\partial x^2} + \frac{\partial^2 \phi}{\partial y^2} + \frac{\partial^2 \phi}{\partial z^2} = 0$

becomes $-4\pi^2 (k_x^2 + k_y^2) \phi(\vec{k}, z) + \frac{\partial^2 \phi(\vec{k}, z)}{\partial z^2} = 0$

General solⁿ is

$$\phi(\vec{k}, z) = A(\vec{k}) e^{-2\pi |k| z} + B(\vec{k}) e^{2\pi |k| z}$$

but $B(\vec{k}) \equiv 0$ since $\phi(\vec{k}, z)$ vanishes @ ∞

$$\phi(\vec{k}, z) = \phi_0(\vec{k}, z_0) = A(\vec{k}) e^{-2\pi |k| z_0}$$

$$|k| = (k_x^2 + k_y^2)^{1/2}$$

$$\downarrow \quad \downarrow$$

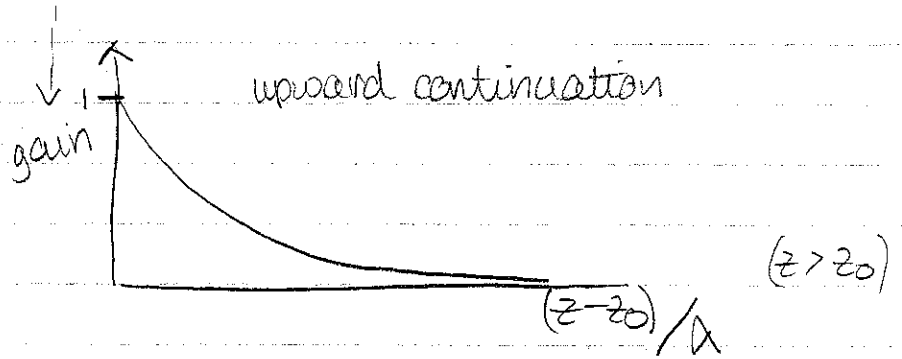
$$\frac{k_x}{2\pi} \quad \frac{k_y}{2\pi}$$

so $\Phi(\vec{r}, z) = \Phi_0(k, z_0) e^{-2\pi|k|(z-z_0)}$

pot' @ any $z > 0$ = pot' @ $z = z_0$ * upward (downward) continuation kernel (factor)
usually measurement altitude

shorthand

$\Phi(z) = \Phi_0 e^{-2\pi \frac{(z-z_0)}{\lambda}}$



Note short λ s \downarrow faster than long λ s

SO If you measure the gravity anomaly @ surface of earth $\Delta g(\vec{x}, 0)$ then can calc @ any height z by

- i) FT $\Delta g(\vec{x}, 0) \rightarrow \Delta g(\vec{k}, 0)$
- ii) multiply by $e^{-2\pi|k|z} \rightarrow \Delta g(\vec{k}, z) = \Delta g(\vec{k}, 0) e^{-2\pi|k|z}$
- iii) Inverse FT result

$\Delta g(x, y, z)$

The upward continuation kernel is WHY it is hard or impossible to recover small-scale features from measurements @ altitude.

From before

$$u(x, y, z) = \frac{1}{2\pi} \iint e^{i(ux+vy)} F(u, v) e^{-(u^2+v^2)^{1/2} z} du dv \quad \text{--- (1)}$$

$$u(x, y, 0) = \frac{1}{2\pi} \iint e^{i(ux+vy)} F(u, v) du dv \quad \text{--- (2)}$$

$$(2) \Rightarrow F(u, v) = \mathcal{F}[u(x, y, 0)] = u\left(\frac{x}{\lambda}, 0\right)$$

$$(1) \Rightarrow F(u, v) e^{-(u^2+v^2)^{1/2} z} = \mathcal{F}[u(x, y, z)] = u\left(\frac{x}{\lambda}, z\right)$$

$$\Rightarrow u\left(\frac{x}{\lambda}, z\right) = u\left(\frac{x}{\lambda}, 0\right) e^{-(u^2+v^2)^{1/2} z}$$

Here $\mathcal{Q}(\vec{k}, z) = \mathcal{Q}(\vec{k}, z_0) e^{-2\pi|\vec{k}|(z-z_0)}$