Chapter 8: A1e, A4e, A7e, A8e Chapter 11: A11e, A22e, A23e Chapter 12: A1e, A7e, A8e, A11e, E21 Total marks out of 54

Chapter 8

A1e)
(2.5 marks)

Using Fig. 8.4, identify whether the following (μ m) are in a window, dirty window, shoulder, or opaque part of the transmittance spectrum, and identify which sketch in Fig. 8.2 shows how the Earth would look at that wavelength. [Hint: transmittance of \geq 80% indicates a window.] e) 1.33

Solution:

1.33 μ m is at a window.

This wavelength has high transmittance so would look like Fig 8.2 a.

<u>Discussion</u>: A 1.33 μ m wavelength is in the infrared spectrum. There is no molecular emittance and very little molecular scattering at this wavelength, leading to transmittance of ~80%. This means that the satellite sees the earth and clouds quite clearly.

A4e)

(2.5 marks)

Find the brightness temperature for the following wavelengths (µm), given a radiance of 10^(-15) W/(m^2 * µm *sr): e) 4.4

ТВ =	71.5: -201.64	1 К 4 °С	
where:	c1B = c2 =	1.: 1.4	.19E+08 W*m-2*μm4*sr-1 .44E+04 μm*K
Use eqn. a	8.2: $T_B = -$	ln(1+	$\frac{c_2 / \lambda}{+ \frac{c_{1B} \cdot \lambda^{-5}}{B_{\lambda}}}$
Find:	TB =	?	К
Given:	λ (μm) = Βλ =		4.4 1E-15 W/(m^2 * μm * sr)

Check: Units ok. Physics ok.

Discussion: The radiance given in this question is very small so the resulting brightness temperature is also small. If the wavelength was smaller, we would get a larger brightness temperature because shorter wavelengths carry more energy.

A7e) (3 marks)	For the following altitudes (km) above the Earth's surface, find the satel orbital periods: e) 5,000.		
	Given:	satellite altitude =	5000 km 5000000 m
	Find:	t_orbit = ?	S
	Use eqn. 8.	8: $t_{orbit} = \frac{2\pi \cdot R^3}{\sqrt{G} \cdot N}$	/2 1
	where:	G = 6.67E-1 M = 5.97E+2	1 N*m^2/kg^2 4 kg
	r_Earth =	6378 km	6378000 m
	R = radius c	of Earth + satellite altitud	e: 11378000 m
	t_orbit =	12076.45 s 3.35 hr	
	Check: Unit	s ok. Physics ok.	
	Discussion: they need le	Satellites at higher altitu ess energy to stay in orbit	des move slower because

A8e) What shade of grey would the following clouds appear in visible, IR, and (3.5 marks) water-vapor satellite images? e) altostratus

Given: altostratus cloud

Find:	Shade of grey in visible, IR, and
	water-vapor satellite images.

VIS:	White during the day.
IR:	Medium gray becase mid-altitude and medium temperature
WV:	Medium grey because altostratus clouds carry some moisture

Discussion: The different shades of grey in different satellite images is indicative of the different wavelengths being picked up by that satellite channel.

Chapter 11

A11e)	Find the ma	gnitude of th	e therma	al wir	nd (m/s) for the	e following t	hickness
(5.5 marks)	gradients: e) ΔTH(km)/Δ	x(km) =	-0.2/	600, ΔTH(km)/	′Δy(km) = -0	.1/400.
	Given:	$\Delta TH_x =$		-0.2	km	-200	m
		∆x =		600	km	600000	m
		∆TH_y =		-0.1	km	-100	m
		Δy =		400	km	400000	m
	Find:	MTH =	?		m/s		
	First, use eq	ns 11.15a an	d 11.15b	:			
					8 ATH		
	($u_{TH} = u_{GS}$	$_{2} - u_{G1}$	=	$\frac{1}{f_c} \Delta y$		
	1	$V_{TH} = V_{G2}$	$-V_{G1}$	= +	$ g \Delta TH$		
					$f_c \Delta x$		
	where lal -		m/s^2				
	assume fc -	1 10F-04	/c				
		1.100 04	73				
	<u> IITH =</u>	22.27	m/s				
	VTН =	-29 70	m/s				
	• • • • •	23170	, s				
	Then use ec	ın 11.16 [.]	Marro	= .	$I_{mx^{2}} + V_{mx^{2}}$		
		1	INITH	- 10	TH TH		
	MTH =	37.12	m/s				

Check: Units ok. Physics ok.

Discussion: The thermal wind speed is the difference in the geostrophic wind

speeds at the top and bottom of the layer. Thickness (TH) is analagous to the mean temperature of the layer. From the given thickness gradients, we know that the temperature is colder to the east, and to the north. Thermal wind points parallel to the thickness isolines, with the cold air to its left. This matches with our numerical answers for UTH and VTH.



Check: Units ok. Physics ok.

Discussion: When we assume that the Coriolis parameter fc changes linearly with latitude (i.e. beta = constant), we are using the "beta plane". This approximation is valid when we are only looking at a small latitude belt.

A23e) (6 marks)	Suppose the average wind speed is 60 m/s from the west at the tropopause For a barotropic Rossby wave at 50° latitude, find both the intrinsic phase speed (m/s) and the phase speed (m/s) relative to the ground for wavelength (km) of: e) 3000.				
	Given:	λ = φ =		3000 km 50 °	3000000 m
	Find:	co = c =	? ?	m/s m/s	

First, find the beta parameter using eqn. 11.35:

$$\beta = \frac{\Delta f_c}{\Delta y} = \frac{2 \cdot \Omega}{R_{earth}} \cdot \cos \phi$$

where :	$2^*\Omega/\text{Rearth} =$	2.29E-11 1/(m*s)
β =	1.47E-11 1/(m*s)	

To find the intrinsic phase speed use eqn. 11.37:

$$c_o = -\beta \cdot \left(\frac{\lambda}{2\pi}\right)^2$$

co =	-3.36 m/s

To find the phase speed use eqn 11.38:

$$c = U_o + c_o$$

Check: Units ok. Physics ok.

Discussion: The intrinsic phase speed tells us how fast the wave is moving relative to the mean flow. A negative intrinsic phase speed means that the waves are actually moving westwards if there is no mean flow in the background. The phase speed is the speed of the wave relative to the ground. This wave travels fairly fast (near mean flow wind speed), which is reasonable as it has a fairly short wavelength for a barotropic wave.

Chapter 12

Identify typical characteristics of the following airmass: e) cE.

(2.5 marks)

A1e)

Given: Airmass = cE = continental equatorial.

Characteristics: Hot and dry. Formed over land.

Discussion:

A continental equatorial airmass is not an airmass we talk about in North America. Forming over equatorial South America or Africa, these airmasses are very hot and dry.

A7e)Find the external Rossby radius of deformation at 60° latitude for a cold(4.5 marks)airmass of thickness 500m and Δθ (°C) of: e) 10.Assume a background temperature of 300K.

Given:	Δθ (°C) = φ (°) = Η (m) = Τ (K) =	10 or 10K because it's a difference 60 500 300			
Find:	λR = ?	m			
Use eqn. 12.5: $a = \lambda_R = \frac{\sqrt{ g \cdot H \cdot \Delta \Theta_v / \overline{T_v}}}{f_c}$					
where g = 9.8 m/s^2					
To find fc use eqn 10.16: fc = $2*\Omega*\sin\phi$ where $\Omega = 7.29E-05$ /s					
fc =	1.26E-04 /s				
Assume a d	ry air mass so that Δ	$\theta \sim = \Delta \theta v$, and T $\sim =$ Tv.			

λR (m) =	1.01E+05 m
	101.22 km

Check: Units ok. Physics ok.

Discussion: above the external Rossby radius of deformation is where the jet

stream would be. The fact that cold air masses cannot redistribute cold air to the equator means some other forces will try to redistribute this heat.

A8) Find and plot the airmass depth and geostrophic wind as a function of (10.5 marks) distance from the front for the cases of the previous exercise. Assume a background potential temperature of 300K.

Given:	Δθ (°C) =	10
	φ (°) =	60
	H (m) =	500
	T (K) =	300

From A7d the final spillage distance of the front a = λR = 101.22 km. a (m) = 1.01E+05

Use eqn 12.6:

$$U_g = -\sqrt{|g| \cdot H \cdot (\Delta \Theta_v / \overline{T_v})} \cdot \exp\left(-\frac{y+a}{a}\right)$$

where y is the distance behind a. and g = 9.8 m/s^2

Assume a dry air mass so that $\Delta \theta \simeq \Delta \theta v$, and T $\simeq Tv$.

Use eqn 12.7:

$$h = H \cdot \left[1 - \exp\left(-\frac{y+a}{a}\right) \right]$$

y (km)		Ug (m/s)	h (km)
	-90	-11.44	0.0524
	-80	-10.36	0.0946
	-70	-9.39	0.1327
	-60	-8.51	0.1672
	-50	-7.71	0.1986
	-25	-6.02	0.2645
	0	-4.70	0.3161
	25	-3.67	0.3563

50	-2.87	0.3878
75	-2.24	0.4123
100	-1.75	0.4315
125	-1.37	0.4465
150	-1.07	0.4582
175	-0.83	0.4674
200	-0.65	0.4745
225	-0.51	0.4801
250	-0.40	0.4844
275	-0.31	0.4878
300	-0.24	0.4905
325	-0.19	0.4926
350	-0.15	0.4942
375	-0.12	0.4955
400	-0.09	0.4965
425	-0.07	0.4972
450	-0.06	0.4978
475	-0.04	0.4983
500	-0.03	0.4987
525	-0.03	0.4990
550	-0.02	0.4992
575	-0.02	0.4994

Check: Units ok. Physics ok.

Discussion: this h vs y plot confirms Fig. 12.17b, where the depth of the cold air mass increases with distance into the cold air mass. Fig 12.17b also shows Ug being strongest at the air mass boundary (y=-a) and then decreasing in strength further back into the cold air mass. Negative values indicate Ug is blowing to the West.



A11e)

(7 marks)

Plot dryline movement with time, given the following conditions. Surface heat flux is constant with time at kinematic rate 0.2 K*m/s. The vertical gradient of potential temperature in the initial sounding is Υ . Terrain slope is s = $\Delta z/\Delta x$. e) Υ (K/km) = 12, s = 1/500.

Given:	Υ =	12	<th>0.012 K/m</th>	0.012 K/m
	s =	1/500 =	0.002	
	FH =	0.2 H	<*m/s	

Plot Δx vs Δt (the dryline movement with time).

From Sample Application, QAK = $FH^*\Delta t$.

Use eqn 12.15:

$$\Delta x = \frac{1}{s} \cdot \left(\frac{2 \cdot Q_{Ak}}{\gamma}\right)^{1/2}$$



12.5	612.37
13	624.50
13.5	636.40
14	648.07
14.5	659.55
15	670.82
15.5	681.91
16	692.82

Check: Units ok. Physics ok.

Discussion: This plot does not take into account night time, when convective turbulence ceases and prevailing low altitude easterlies advect moist air back towards the west.

E21)

(4 marks) Background: Recall that a frontal zone separates warmer and cooler airmasses. The warm airmass side of this zone is where the front is drawn on a weather map. This is true for both cold and warm fronts. Issue: AFTER passage of the cold front is when significant temperature decreases are observed. BEFORE passage of a warm front is when significant warming is observed. Question: Why does this difference exist (ie. AFTER vs BEFORE) for the passage of these two fronts?

Solution:

Fronts on a weather map are always drawn on the warm side of the surface frontal zone.

For a warm front, the baroclinic zone (the region with the strongest horizontal temperature gradient indicated by the tightest isotherms) is ahead of the warm front. For this reason, you will feel the rapid warming before the warm front approaches.

For a cold front the baroclinic zone is **behind** the cold front, so you will feel the rapid cooling after the cold front has passed.