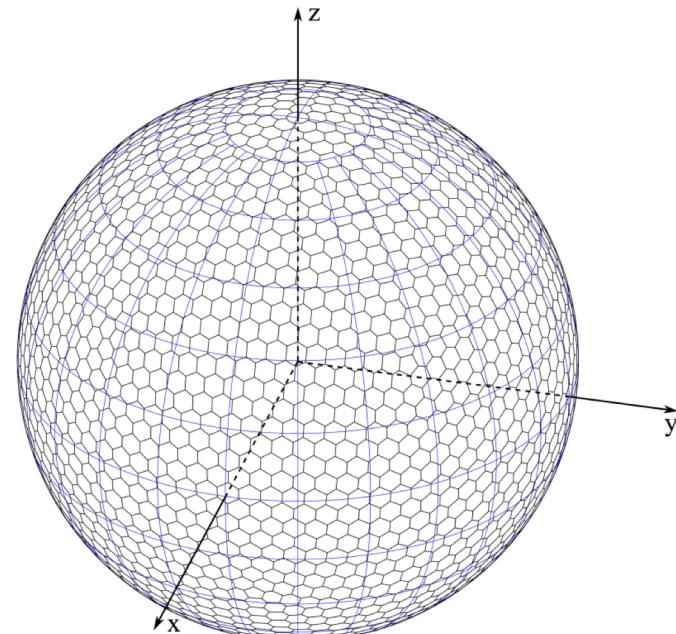


MPAS NUMERICS, DYNAMICS, AND RESEARCH

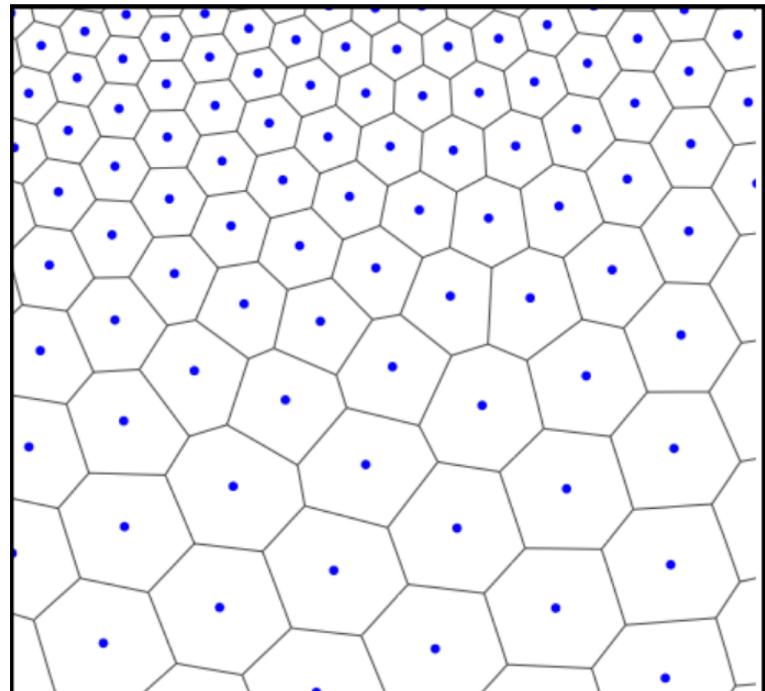
Timothy Chui

Mar. 21, 2019



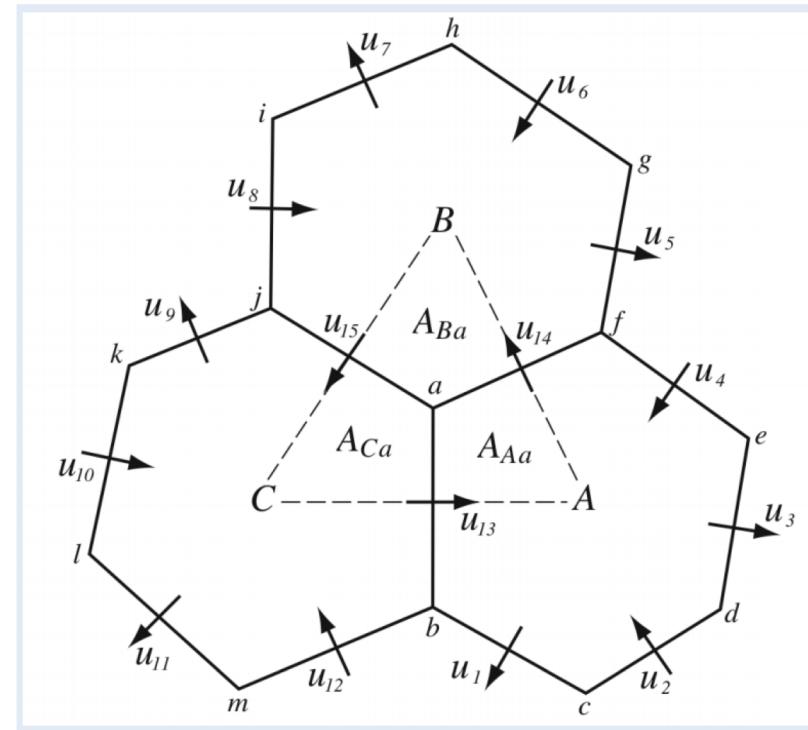
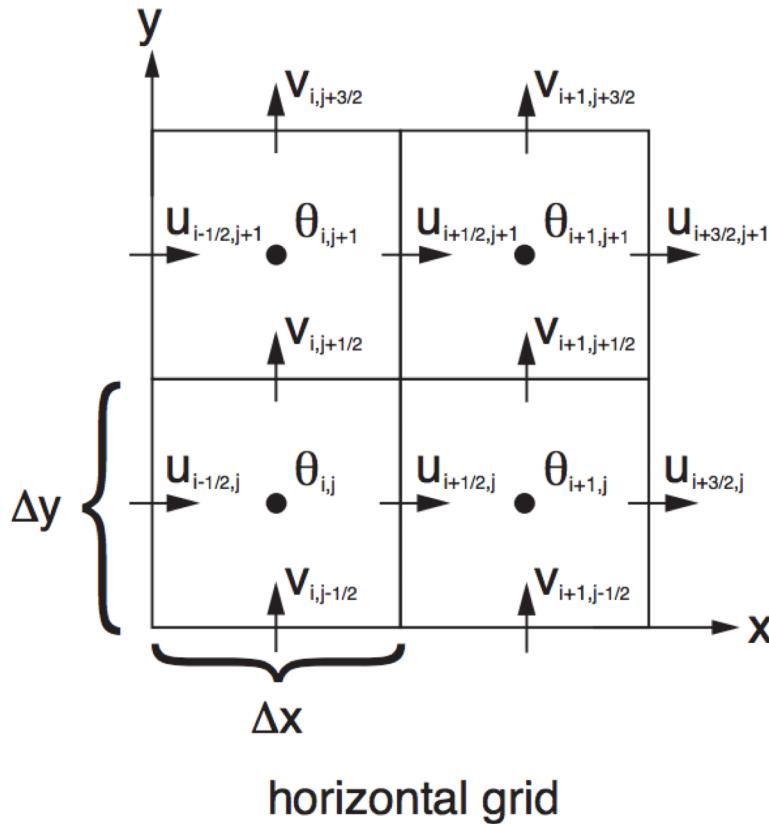
Overview of seminar series

- Introduction to MPAS
- The MPAS Mesh
- Introduction to Finite-Volume Discretization
- **MPAS Numerics, Dynamic, and Research**



<http://mpas-dev.github.io/atmosphere/tutorial.html>

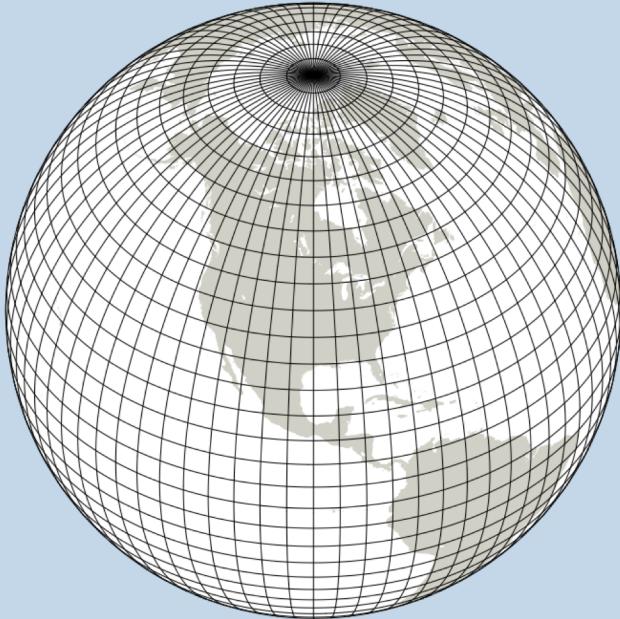
Seminar 1: WRF vs. MPAS



<http://mpas-dev.github.io/atmosphere/tutorial.html>

- Eulerian, finite-difference, **structured**
- Eulerian, finite-volume, **unstructured**

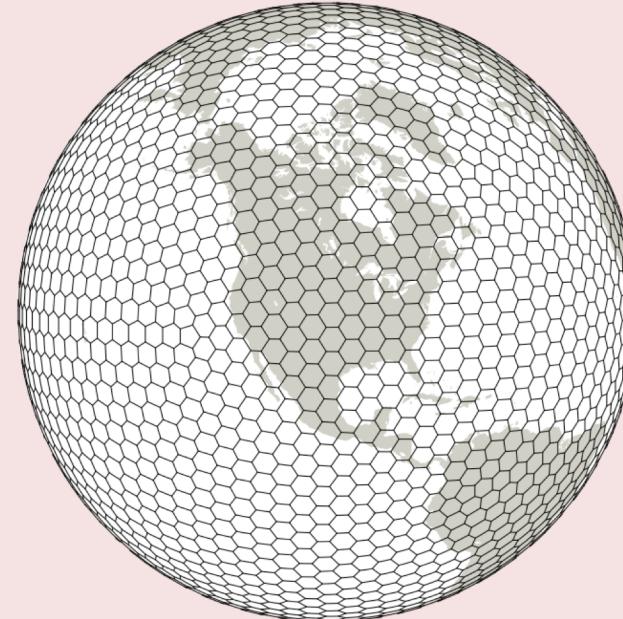
Seminar 1: WRF vs. MPAS



WRF

Lat-Lon global grid

- Anisotropic grid cells
- Polar filtering required
- Poor scaling on massively parallel computers



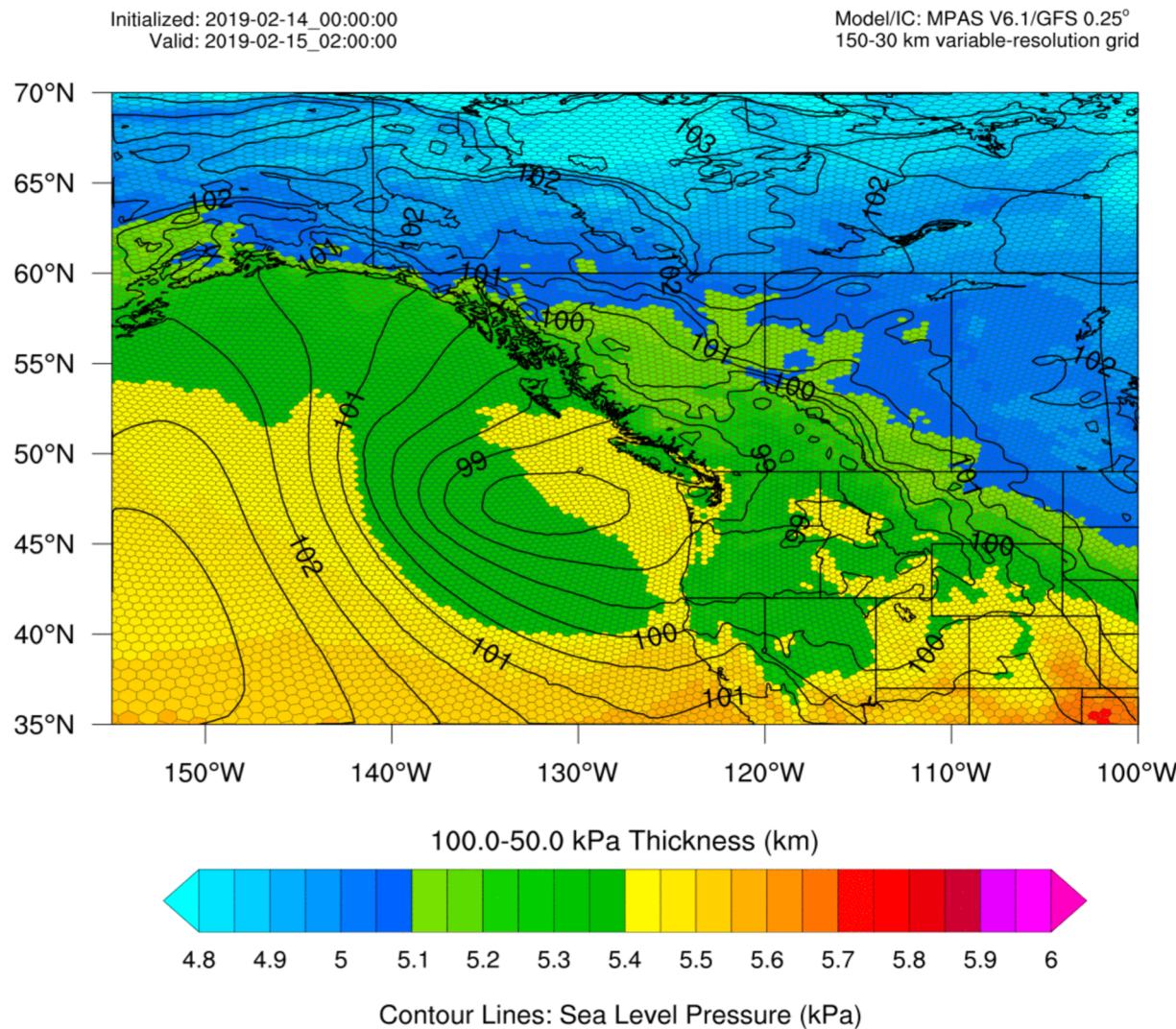
MPAS

Unstructured Voronoi (hexagonal) grid

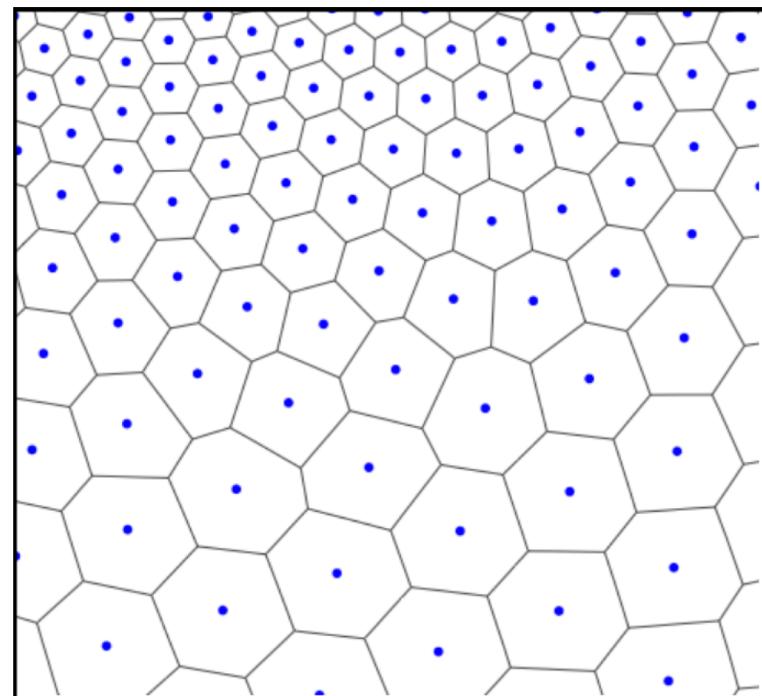
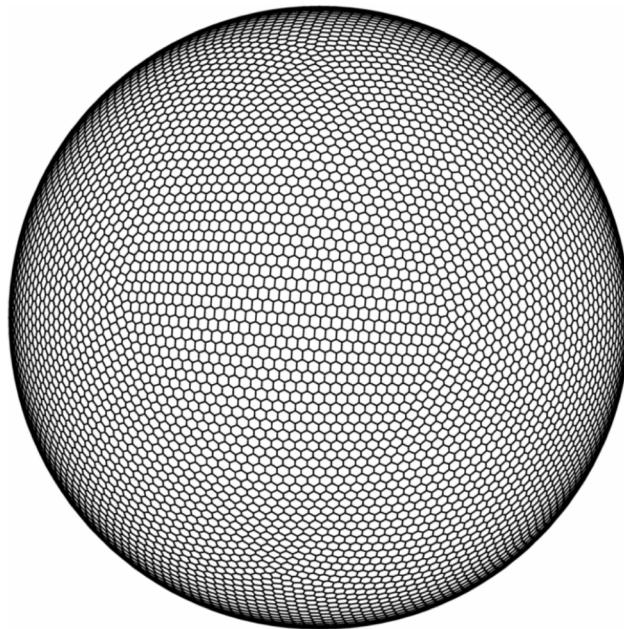
- Good scaling on massively parallel computers
- No pole problems

Seminar 1: MPAS Forecasts

Sea Level Pressure and 100.0-50.0 kPa Thickness



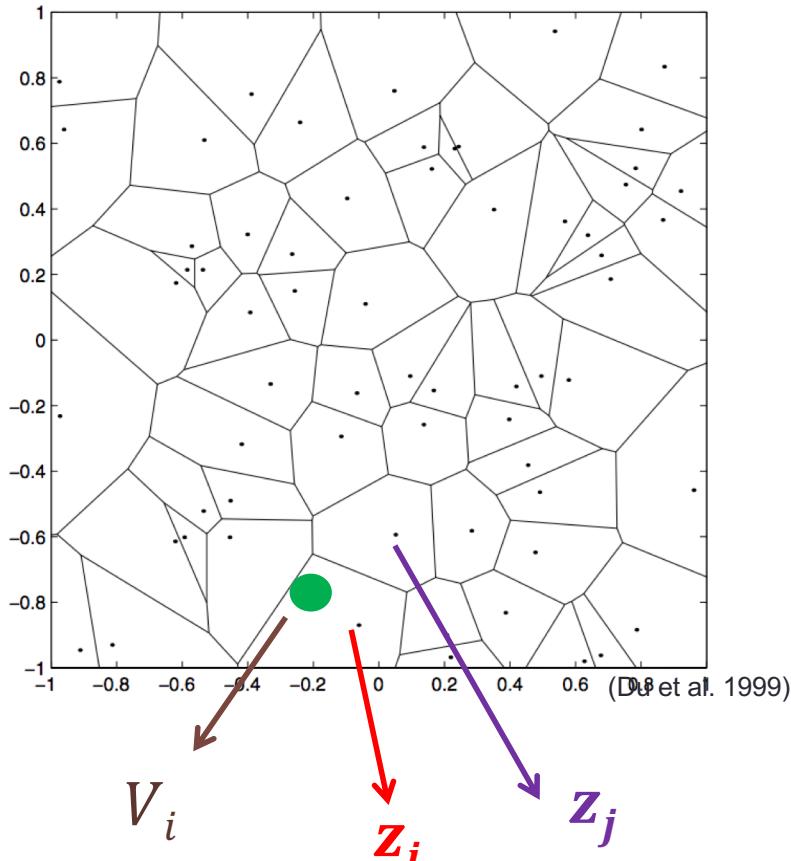
Seminar 2: Spherical Centroidal Voronoi Tessellations



<http://mpas-dev.github.io/atmosphere/tutorial.html>

Seminar 2: Spherical Centroidal Voronoi Tessellations

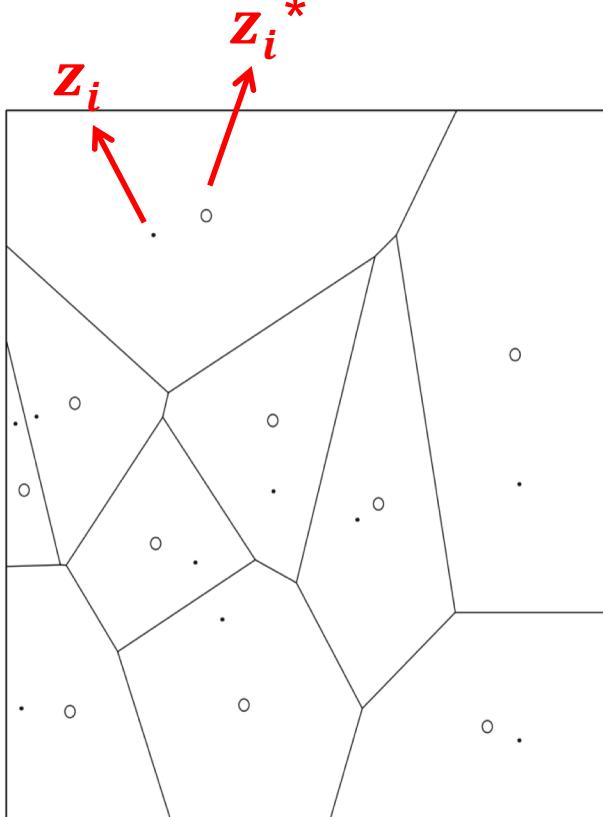
Ω



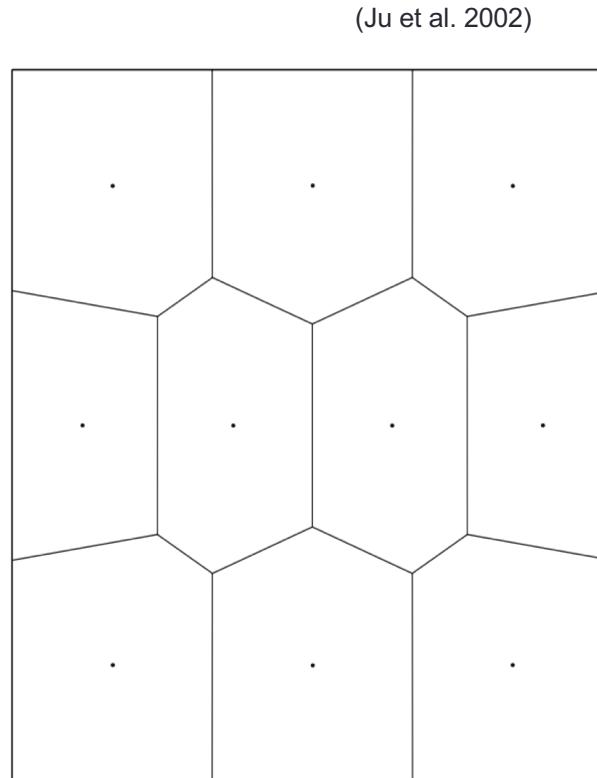
- Points associated with a Voronoi region V_i make up a Voronoi set \widehat{V}_i

$$\widehat{V}_i = \{ \mathbf{x} \in \Omega \mid |\mathbf{x} - \mathbf{z}_i| < |\mathbf{x} - \mathbf{z}_j| \text{ for } j = 1, \dots, k, j \neq i \}$$

Seminar 2: Spherical Centroidal Voronoi Tessellations



$$z_i \neq z_i^*, i = 1, \dots, k$$



$$z_i = z_i^*, i = 1, \dots, k$$

(Ju et al. 2002)

- In general, the generating point z_i of each Voronoi region is not the same as the mass centroid z_i^* of the region

Seminar 3: Finite-Volume Discretization

$$\frac{d\bar{U}}{dt} = -\frac{1}{V} \iint_{\partial CV} (\vec{F} \cdot \hat{n}) dA + \bar{S}$$

- Change in the control volume average \bar{U} is due to stuff leaving and entering control volume, and stuff \bar{S} being created or destroyed on the inside

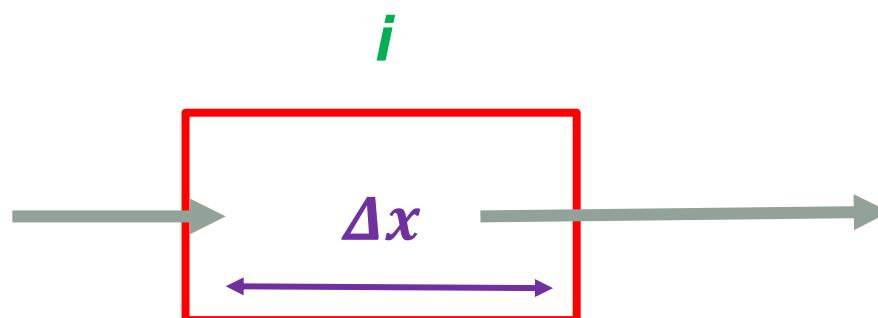
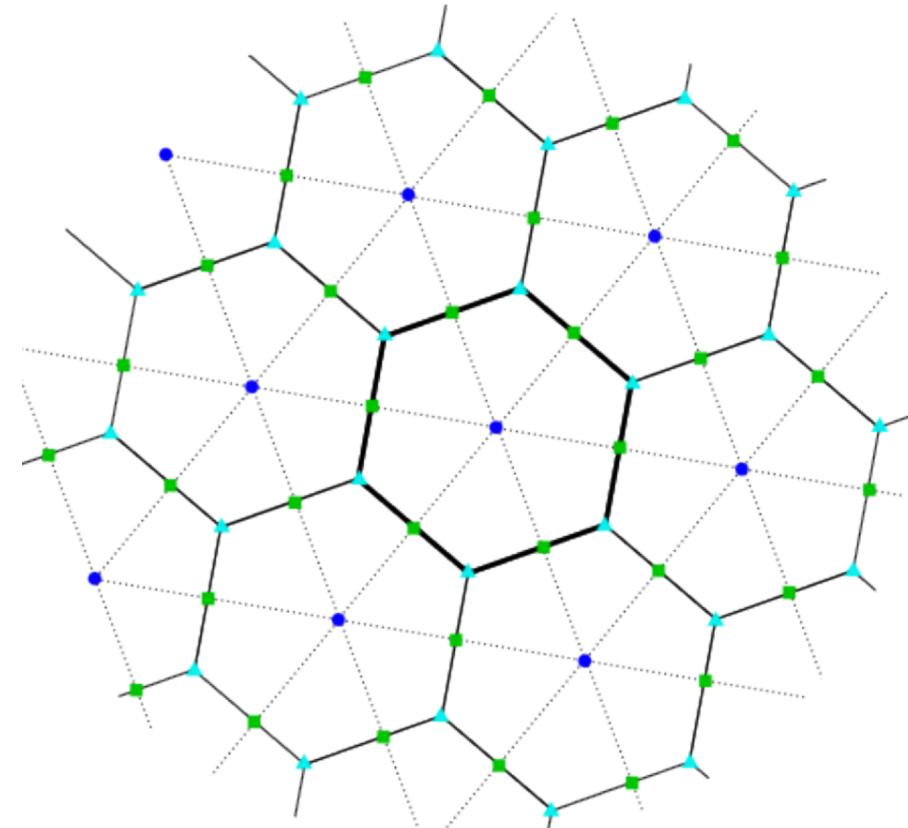


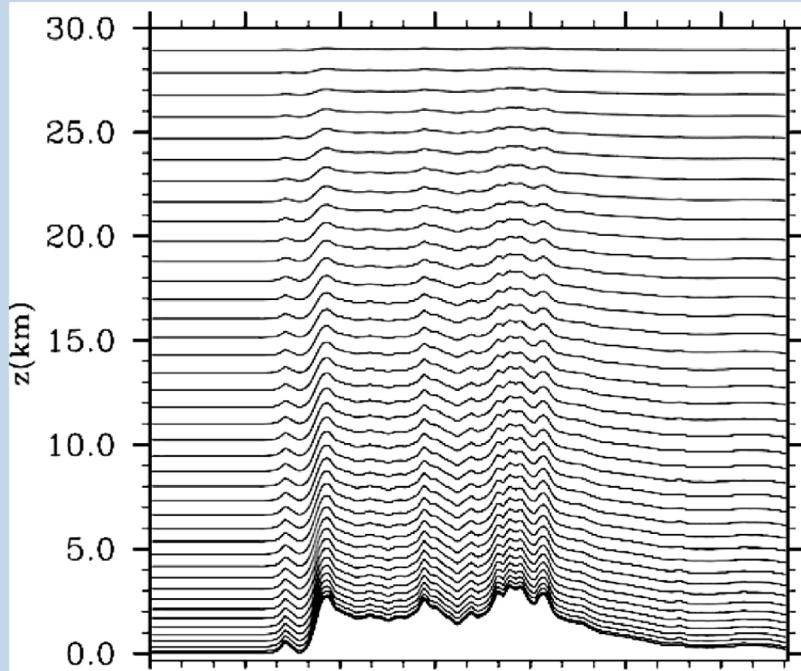
Table of Contents

- Vertical Coordinate System
- Horizontal Coordinate System
- MPAS Dynamical Equations
- Vertical Discretization
- Horizontal Discretization
- Summary

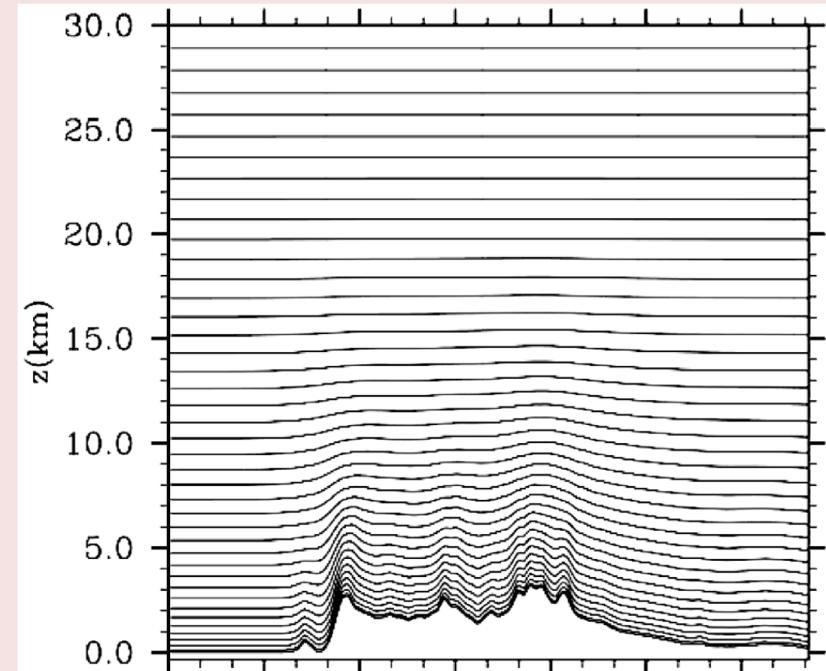


<http://mpas-dev.github.io/atmosphere/tutorial.html>

Vertical Coordinate System



WRF
Pressure-based
terrain-following sigma
vertical coordinate



MPAS
Height-based hybrid smoothed
terrain-following vertical coordinate

- Improved numerical accuracy

Vertical Coordinate System

Basic terrain-following (**BTF**; Gal-Chen and Somerville 1975):

$$z(x, y, \zeta) = \zeta + \left[1 - \frac{\zeta}{z_t}\right] h(x, y)$$

Smoothed terrain-following (**STF**; Klemp 2011)

$$z(x, y, \zeta) = \zeta + A(\zeta)h_s(x, y, \zeta)$$

$z(x, y, \zeta)$ = geometric height [m]

ζ = terrain-following constant; for BTF $\zeta = z$ at z_t (model lid)

$h(x, y)$ = terrain height [m]

$h_s(x, y, \zeta)$ = terrain influence in STF; $h_s(x, y, 0) = h$

$A(\zeta)$ = controls how quickly terrain-following \rightarrow constant height

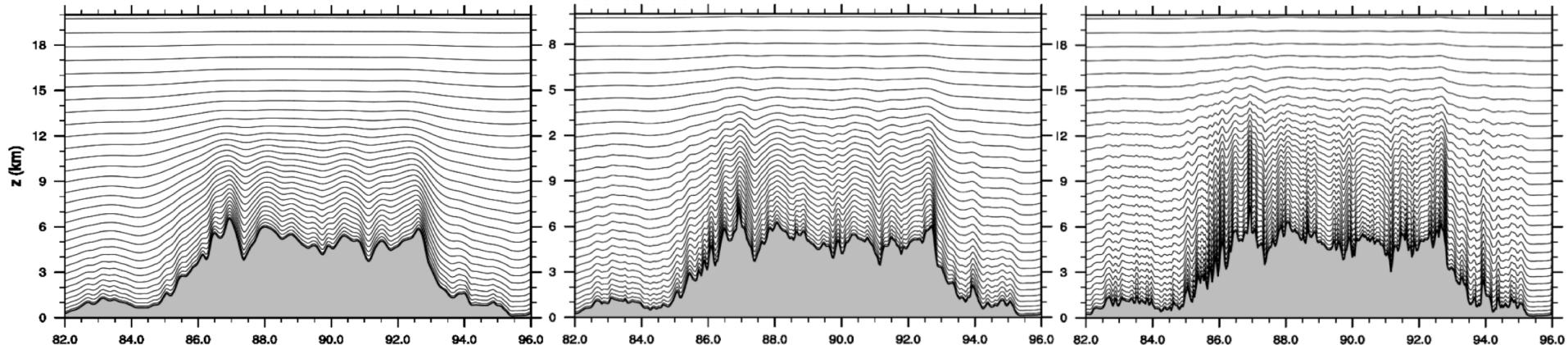
Vertical Coordinate System

15 km grid

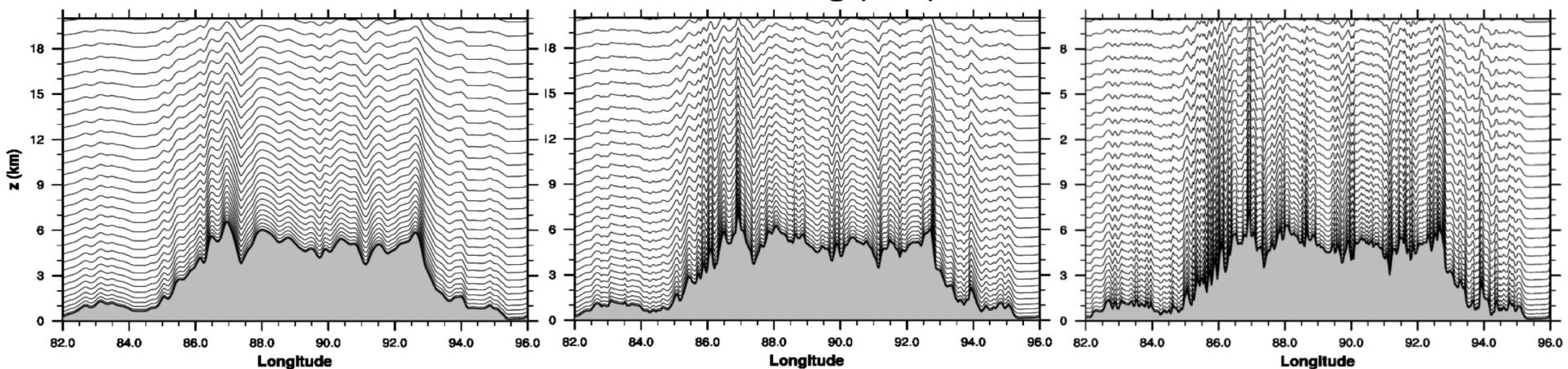
7.5 km grid

3 km grid

Smoothed hybrid terrain-following (STF) coordinate



Basic terrain-following (BTF) coordinate



(Model top is at 30 km)

Vertical Coordinate System - Init

- **Terrain**
 - 30 arcsec data averaged over grid-cell areas
 - Single pass of smoothing using 4th-order Laplacian
- **Vertical Coordinate**
 - Choice non-smoothed (BTF) or n -times smoothed (STF); iterate upwards from $h_s^n(x, y, 0) = h$

$\beta(\zeta)$ = smoothing coefficient

d = grid-cell length scale

z_H = height where $\zeta = z$

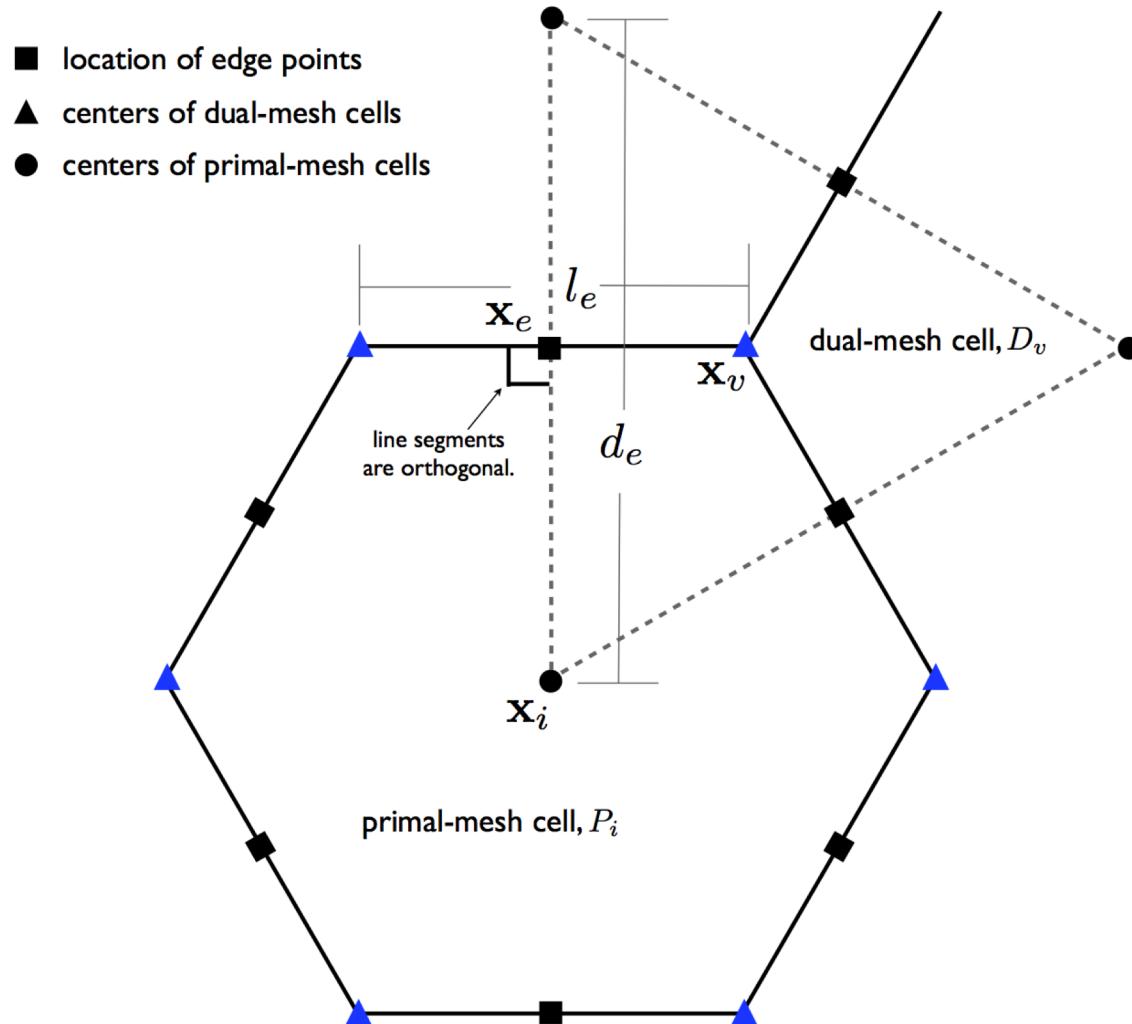
$\nabla_\zeta^2 h_s^n$ = Laplacian

$$h_s^{n+1} = h_s^n + \beta(\zeta) d^2 \nabla_\zeta^2 h_s^n$$

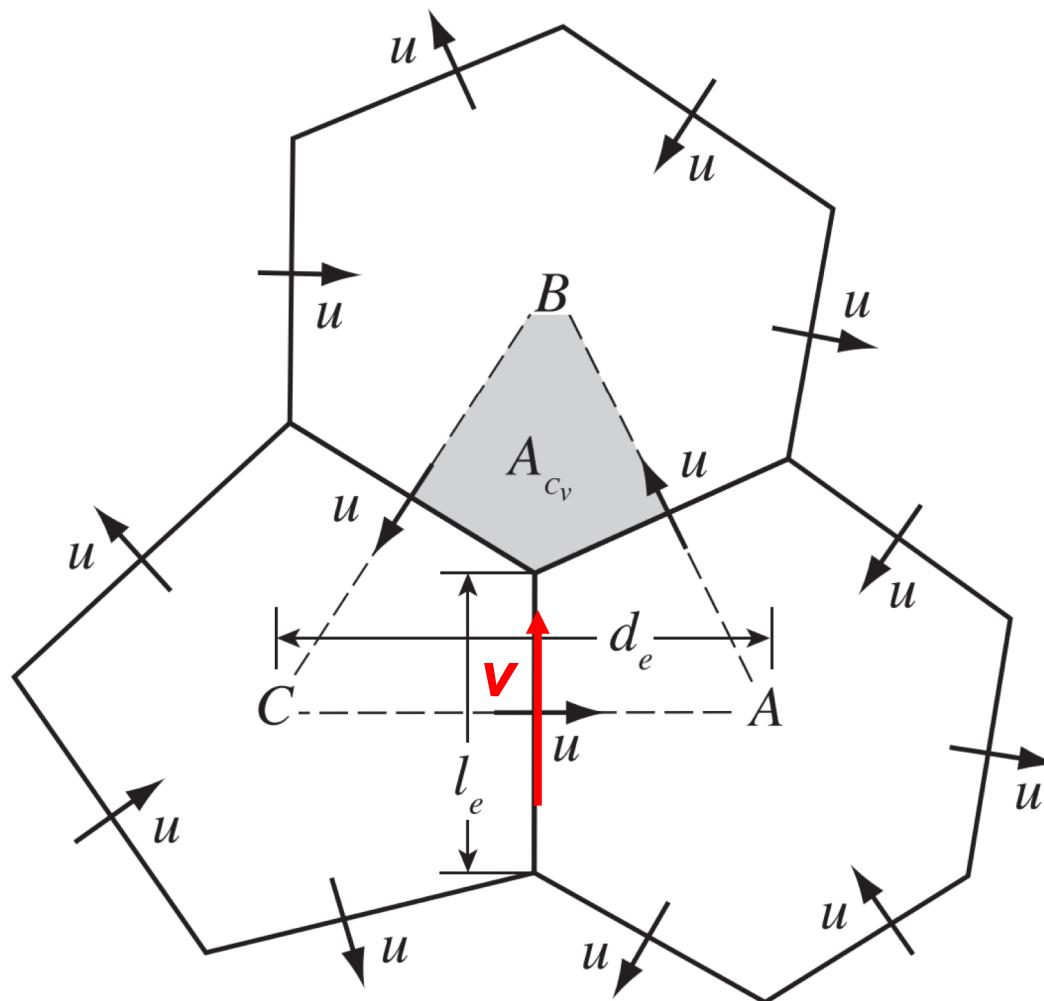
$$A(\zeta) = \cos^6 \left(\frac{\pi}{2} \frac{\zeta}{z_H} \right), \quad \zeta < z_H$$

$$A(\zeta) = 0, \quad \zeta \geq z_H$$

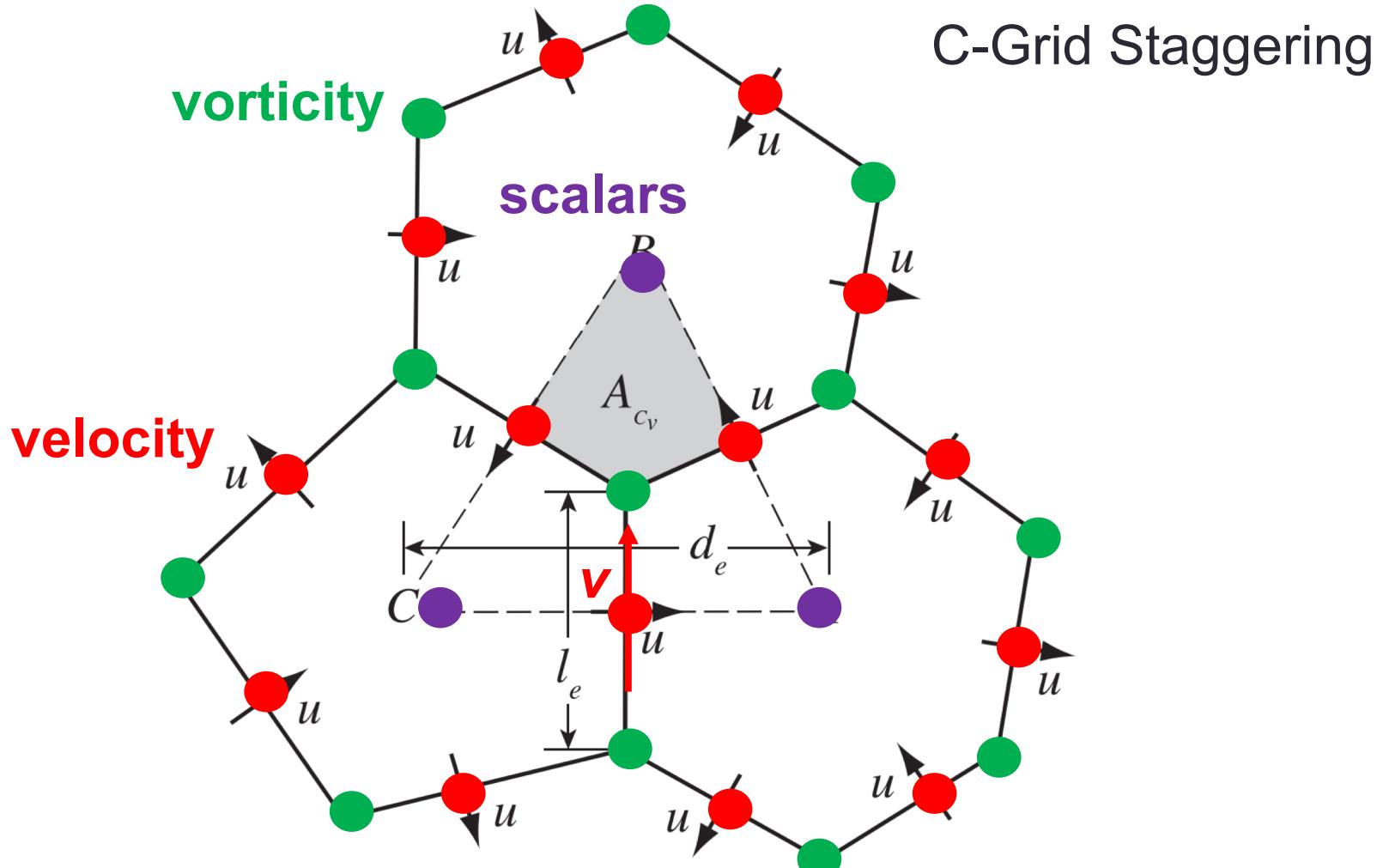
Horizontal Coordinate System



Horizontal Coordinate System



Horizontal Coordinate System



MPAS Dynamical Equations

- Geostrophic flow
- Dry
- Incompressible
- Pythagorean

$$u = -\frac{1}{\rho f} \frac{\partial p}{\partial y}$$

$$v = \frac{1}{\rho f} \frac{\partial p}{\partial x}$$

$$a^2 + b^2 = c^2$$

MPAS Dynamical Equations

- Coupled variables in conservative form
- Hybrid-height coordinate
- Nonhydrostatic
- Fully compressible
- Coupled (moist) potential temperature

$$\theta_m = \theta[1 + (R_v/R_d)q_v]$$

$$p = p_0 \left(\frac{R_d \zeta_z \Theta_m}{p_0} \right)^\gamma$$

$$\frac{\rho_m}{\rho_d} = 1 + q_v + q_c + q_r + \dots$$

$$\begin{aligned} \frac{\partial \mathbf{V}_H}{\partial t} &= -\frac{\rho_d}{\rho_m} \left[\nabla_\zeta \left(\frac{p}{\zeta_z} \right) - \frac{\partial \mathbf{z}_H p}{\partial \zeta} \right] - \eta \mathbf{k} \times \mathbf{V}_H \\ &\quad - \mathbf{v}_H \nabla_\zeta \cdot \mathbf{V} - \frac{\partial \Omega \mathbf{v}_H}{\partial \zeta} - \rho_d \nabla_\zeta K \\ &\quad - eW \cos \alpha_r - \frac{\mathbf{v}_H W}{r_e} + \mathbf{F}_{\mathbf{V}_H}, \\ \frac{\partial W}{\partial t} &= -\frac{\rho_d}{\rho_m} \left[\frac{\partial p}{\partial \zeta} + g \tilde{\rho}_m \right] - (\nabla \cdot \mathbf{v} W)_\zeta + \frac{uU + vV}{r_e} \\ &\quad + e(U \cos \alpha_r - V \sin \alpha_r) + F_W, \end{aligned}$$

$$\frac{\partial \Theta_m}{\partial t} = -(\nabla \cdot \mathbf{V} \theta_m)_\zeta + F_{\Theta_m},$$

$$\frac{\partial \tilde{\rho}_d}{\partial t} = -(\nabla \cdot \mathbf{V})_\zeta, \quad \text{and}$$

$$\frac{\partial Q_j}{\partial t} = -(\nabla \cdot \mathbf{V} q_j)_\zeta + F_{Q_j}.$$

MPAS Dynamical Equations

- Coupled variables in conservative form
- Hybrid-height coordinate
- Nonhydrostatic
- Fully compressible
- Coupled (moist) potential temperature

Moist (virtual) θ

$$\theta_m = \theta[1 + (R_v/R_d)q_v]$$

$$p = p_0 \left(\frac{R_d \zeta_z \Theta_m}{p_0} \right)^\gamma \quad \text{Eqn of state}$$

$$\frac{\rho_m}{\rho_d} = 1 + q_v + q_c + q_r + \dots$$

Moist air

$$\frac{\partial \mathbf{V}_H}{\partial t} = -\frac{\rho_d}{\rho_m} \left[\nabla_\zeta \left(\frac{p}{\zeta_z} \right) - \frac{\partial \mathbf{z}_H p}{\partial \zeta} \right] - \eta \mathbf{k} \times \mathbf{V}_H$$

$$- \mathbf{v}_H \nabla_\zeta \cdot \mathbf{V} - \frac{\partial \Omega \mathbf{v}_H}{\partial \zeta} - \rho_d \nabla_\zeta K$$

$$- eW \cos \alpha_r - \frac{\mathbf{v}_H W}{r_e} + \mathbf{F}_{\mathbf{V}_H},$$

$$\frac{\partial W}{\partial t} = -\frac{\rho_d}{\rho_m} \left[\frac{\partial p}{\partial \zeta} + g \tilde{\rho}_m \right] - (\nabla \cdot \mathbf{v} W)_\zeta + \frac{uU + vV}{r_e}$$

$$+ e(U \cos \alpha_r - V \sin \alpha_r) + F_W,$$

$$\frac{\partial \Theta_m}{\partial t} = -(\nabla \cdot \mathbf{V} \theta_m)_\zeta + F_{\Theta_m},$$

$$\frac{\partial \tilde{\rho}_d}{\partial t} = -(\nabla \cdot \mathbf{V})_\zeta, \quad \text{and}$$

$$\frac{\partial Q_j}{\partial t} = -(\nabla \cdot \mathbf{V} q_j)_\zeta + F_{Q_j}.$$

Hor. momentum

Vert. momentum

Temperature
Continuity

Moisture

MPAS Dynamical Equations

- Flux-form prognostic equations

$$(U, V, W, \Theta_m, Q_j) = \widetilde{\rho_d} \cdot (u, v, w, \theta_m, q_j)$$

$j = v$ (vapour), c (cloud), r (rainwater), etc.

$$\widetilde{\rho_d} = \rho_d \left(\frac{1}{\frac{d\zeta}{dz}} \right) = \rho_d \zeta_z$$

$$\boldsymbol{v} = (u, v, w), \boldsymbol{V} = (U, V, W)$$

MPAS Dynamical Equations

- Vertical momentum

$$\begin{aligned}\frac{\partial W}{\partial t} = & -\frac{\rho_d}{\rho_m} \left[\frac{\partial p}{\partial \zeta} + g \tilde{\rho}_m \right] - (\nabla \cdot \mathbf{v}W)_{\zeta} + \frac{uU + vV}{r_e} \\ & + e(U \cos\alpha_r - V \sin\alpha_r) + F_W,\end{aligned}$$

MPAS Dynamical Equations

- Vertical momentum

$$\frac{\partial W}{\partial t} = -\frac{\rho_d}{\rho_m} \left[\frac{\partial p}{\partial \zeta} + g \tilde{\rho}_m \right] - (\nabla \cdot \mathbf{v} W)_{\zeta} + \frac{uU + vV}{r_e} + e(U \cos \alpha_r - V \sin \alpha_r) + F_W,$$

Coriolis Physics

Pressure gradient Gravity Advection Curvature

MPAS Dynamical Equations

- Horizontal momentum

$$\begin{aligned}\frac{\partial \mathbf{V}_H}{\partial t} = & -\frac{\rho_d}{\rho_m} \left[\nabla_\zeta \left(\frac{p}{\zeta_z} \right) - \frac{\partial \mathbf{z}_H p}{\partial \zeta} \right] - \eta \mathbf{k} \times \mathbf{V}_H \\ & - \mathbf{v}_H \nabla_\zeta \cdot \mathbf{V} - \frac{\partial \Omega \mathbf{v}_H}{\partial \zeta} - \rho_d \nabla_\zeta K \\ & - eW \cos \alpha_r - \frac{\mathbf{v}_H W}{r_e} + \mathbf{F}_{\mathbf{V}_H},\end{aligned}$$

MPAS Dynamical Equations

- Horizontal momentum

$$\frac{\partial \mathbf{V}_H}{\partial t} = -\frac{\rho_d}{\rho_m} \left[\nabla_\zeta \left(\frac{p}{\zeta_z} \right) - \frac{\partial \mathbf{z}_H p}{\partial \zeta} \right] - \eta \mathbf{k} \times \mathbf{V}_H$$

Pressure gradient

Nonlinear Coriolis (Vorticity)

Advection

$$- \mathbf{v}_H \nabla_\zeta \cdot \mathbf{V} - \frac{\partial \Omega \mathbf{v}_H}{\partial \zeta} - \rho_d \nabla_\zeta K$$

Kinetic energy gradient

Coriolis

$$- eW \cos \alpha_r - \frac{\mathbf{v}_H W}{r_e} + \mathbf{F}_{\mathbf{V}_H},$$

Curvature

Physics

Transport (Advection)

- ψ = some quantity

$$\frac{\partial \rho\psi}{\partial t} = -\nabla_\zeta \cdot \mathbf{v}_H(\rho\psi) - \frac{\partial W\rho\psi}{\partial z}$$

Scalar Transport

- FV formulation over, integrate over cell

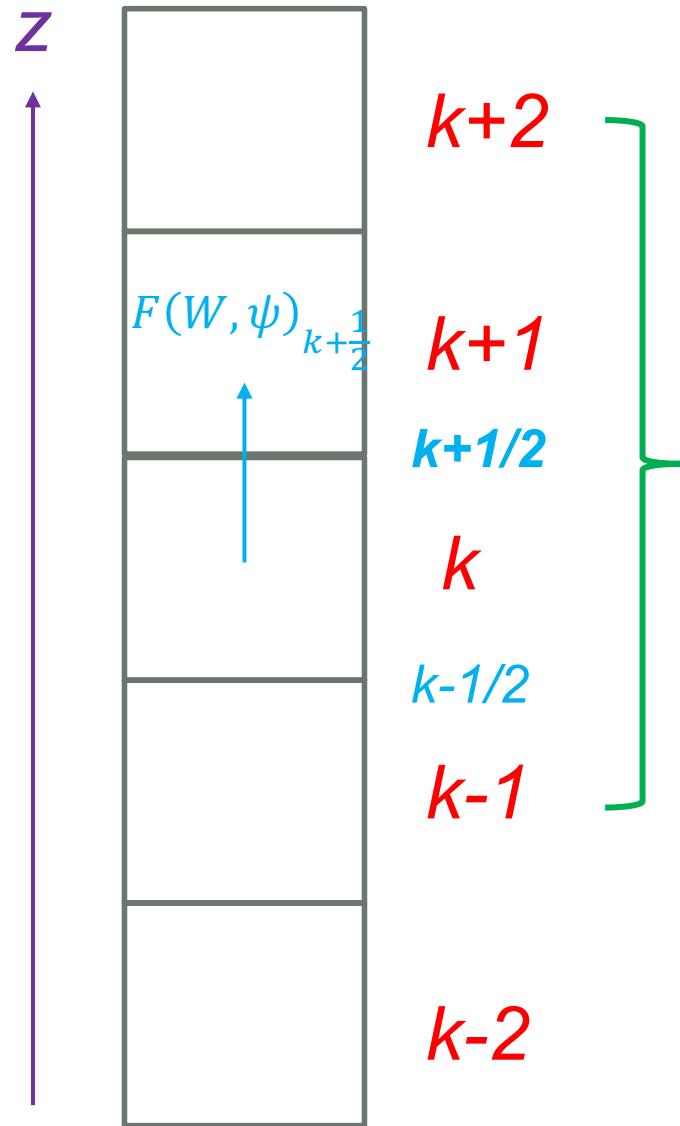
$$\int_V \frac{\partial \rho\psi}{\partial t} dV = - \int_V (\nabla \cdot \mathbf{v}(\rho\psi)) dV$$

Scalar Transport

- Apply divergence theorem; control-volume average definition
- Σ = surface of polyhedron; $d\sigma$ = surface element

$$\frac{\partial \overline{\rho\psi}}{\partial t} = -\frac{1}{V} \int_{\Sigma} (\rho\psi) \mathbf{v} \cdot \mathbf{n} d\sigma$$

Vertical Discretization



$$W = \rho w$$

$$\delta_z^2 \psi_k = \psi_{k-1} - 2\psi_k + \psi_{k+1}$$

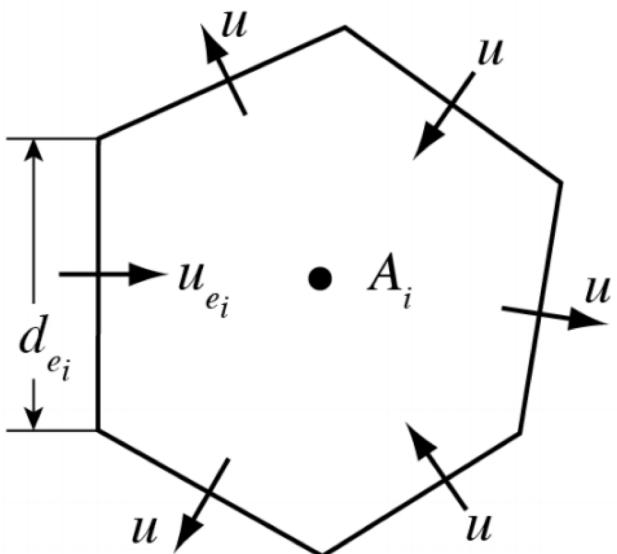
$$\begin{aligned} F(W, \psi)_{k+\frac{1}{2}} &= W_{k+\frac{1}{2}} \left[\frac{1}{2} (\psi_{k+1} + \psi_k) \right. \\ &\quad - \frac{1}{12} (\delta_z^2 \psi_{k+1} + \delta_z^2 \psi_k) \\ &\quad \left. + sign(W) \frac{\beta}{12} (\delta_z^2 \psi_{k+1} - \delta_z^2 \psi_k) \right] \end{aligned}$$

$\beta = 0$: 4th-order scheme, neutral

$\beta > 0$: 3rd -order scheme, damping

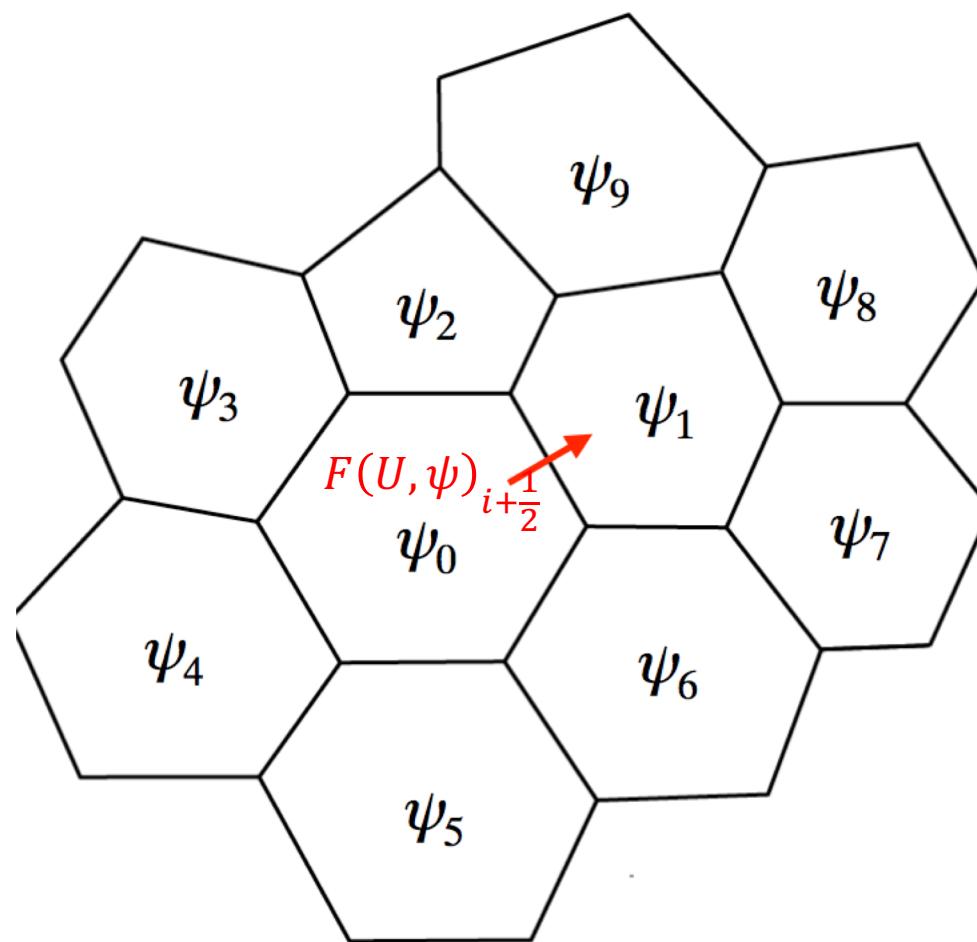
Horizontal Discretization

- Sum counter-clockwise around hexagon edges



$$\begin{aligned} & -\frac{1}{V} \int_{\Sigma} (\rho \psi) \mathbf{v}_H \cdot \mathbf{n} d\sigma \\ & \approx -\frac{1}{A_i} \sum_{n_{e_i}} d_{e_i} \overline{(\rho \mathbf{v}_H \cdot \mathbf{n}_{e_i})} \psi_i \end{aligned}$$

Horizontal Discretization



Horizontal Discretization

$$\begin{aligned} & F(U, \psi)_{i+\frac{1}{2}} \\ &= U_{i+\frac{1}{2}} \left[\frac{1}{2} (\psi_{i+1} + \psi_i) - \frac{1}{12} (\delta_x^2 \psi_{i+1} + \delta_x^2 \psi_i) \right. \\ &\quad \left. + sign(U) \frac{\beta}{12} (\delta_x^2 \psi_{i+1} - \delta_x^2 \psi_i) \right] \end{aligned}$$

- x = direction normal to cell edges
- Coordinates aren't next to each other in the horizontal; can't do straight-forward computation of $\delta_x^2 \psi_i$

$$\delta_x^2 \psi_i = \Delta x^2 \frac{\partial^2 \psi}{\partial x^2} + \mathcal{O}(\Delta x^4)$$

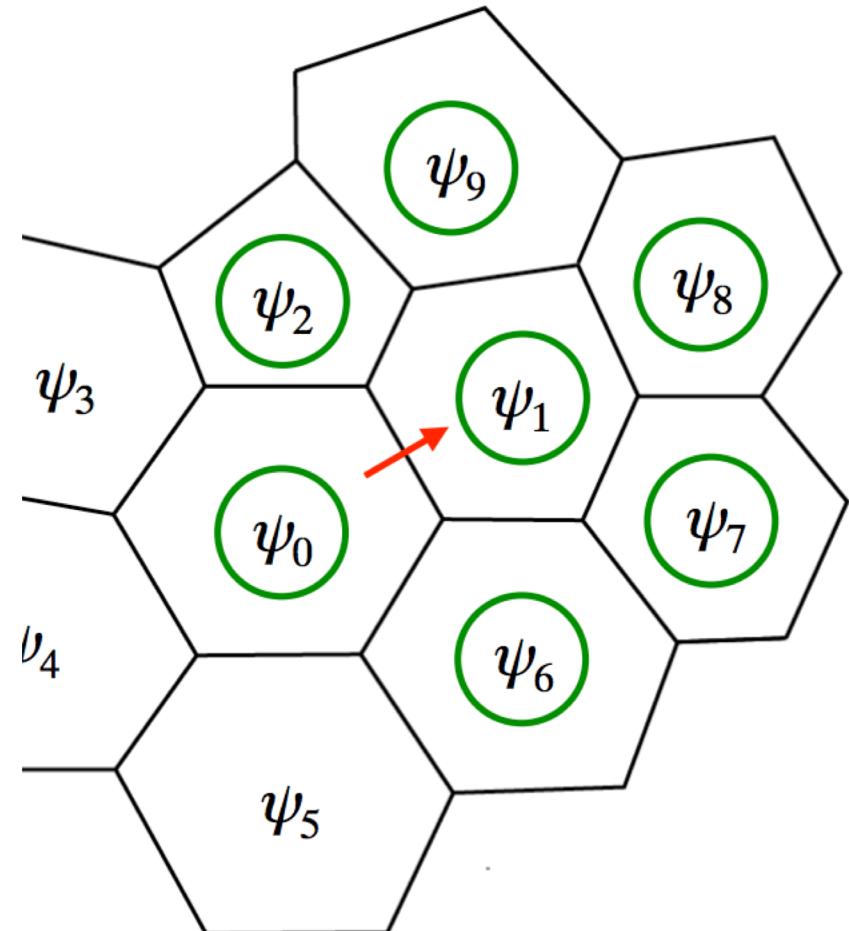
Horizontal Discretization

$$\begin{aligned} & F(U, \psi)_{i+\frac{1}{2}} \\ &= U_{i+\frac{1}{2}} \left[\frac{1}{2} (\psi_{i+1} + \psi_i) \right. \\ &\quad - \Delta x_e^2 \frac{1}{12} \left\{ \left(\frac{\partial^2 \psi}{\partial x^2} \right)_{i+1} + \left(\frac{\partial^2 \psi}{\partial x^2} \right)_i \right\} \\ &\quad \left. + sign(U) \Delta x_e^2 \frac{\beta}{12} \left\{ \left(\frac{\partial^2 \psi}{\partial x^2} \right)_{i+1} - \left(\frac{\partial^2 \psi}{\partial x^2} \right)_i \right\} \right] \end{aligned}$$

Horizontal Discretization

https://www.earthsystemcog.org/site_media/projects/dycore_test_group/20160122_MPAS_configuration_overview.pdf

$$\begin{aligned} F(U, \psi)_{i+\frac{1}{2}} &= U_{i+\frac{1}{2}} \left[\frac{1}{2} (\psi_{i+1} + \psi_i) \right. \\ &\quad - \Delta x_e^2 \frac{1}{12} \left\{ \left(\frac{\partial^2 \psi}{\partial x^2} \right)_{i+1} + \left(\frac{\partial^2 \psi}{\partial x^2} \right)_i \right\} \\ &\quad + sign(U) \Delta x_e^2 \frac{\beta}{12} \left\{ \left(\frac{\partial^2 \psi}{\partial x^2} \right)_{i+1} \right. \\ &\quad \left. - \left(\frac{\partial^2 \psi}{\partial x^2} \right)_i \right\} \end{aligned}$$

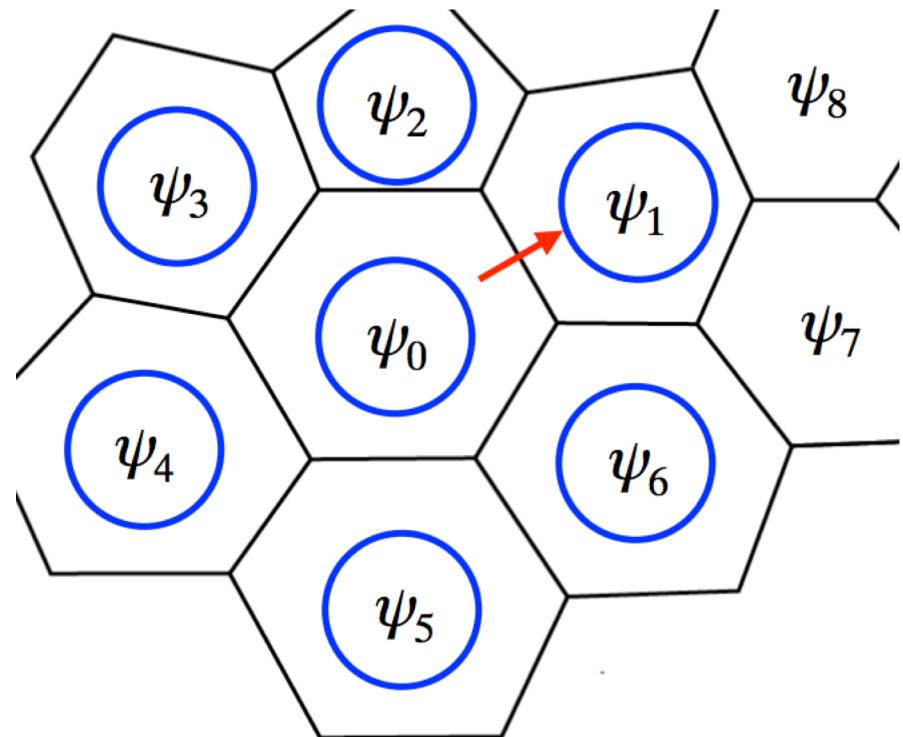


Fit polynomial to approximate second derivatives (Skamarock and Gassmann 2011)

Horizontal Discretization

https://www.earthsystemcog.org/site_media/projects/dycore_test_group/20160122_MPAS_configuration_overview.pdf

$$\begin{aligned}
 & F(U, \psi)_{i+\frac{1}{2}} \\
 &= U_{i+\frac{1}{2}} \left[\frac{1}{2} (\psi_{i+1} + \psi_i) \right. \\
 &\quad - \Delta x_e^2 \frac{1}{12} \left\{ \left(\frac{\partial^2 \psi}{\partial x^2} \right)_{i+1} + \left(\frac{\partial^2 \psi}{\partial x^2} \right)_i \right\} \\
 &\quad + sign(U) \Delta x_e^2 \frac{\beta}{12} \left\{ \left(\frac{\partial^2 \psi}{\partial x^2} \right)_{i+1} \right. \\
 &\quad \left. \left. - \left(\frac{\partial^2 \psi}{\partial x^2} \right)_i \right\}
 \end{aligned}$$



Fit polynomial to approximate second derivatives (Skamarock and Gassmann 2011)

Not discussed

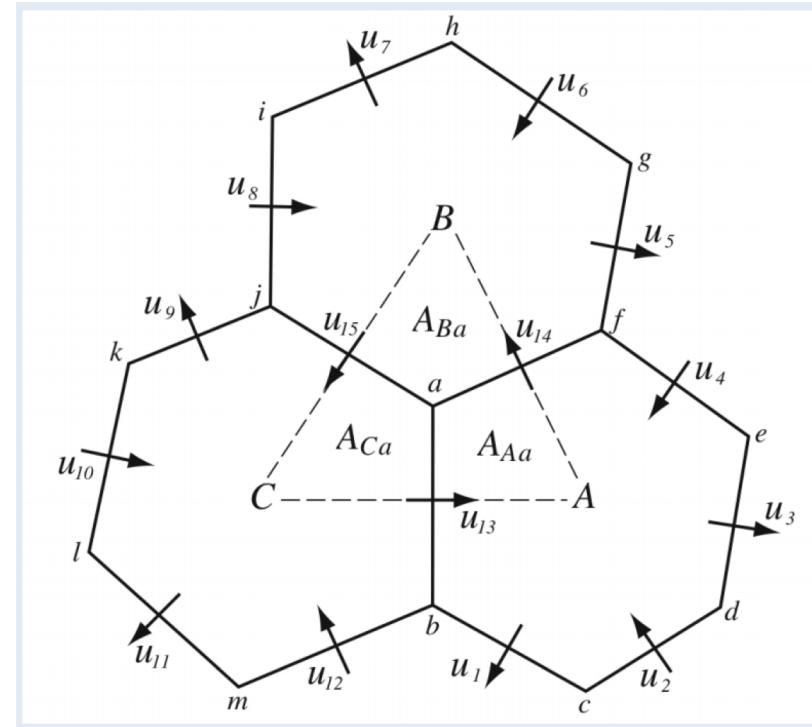


<https://www.convertwithcontent.com/web-marketing-in-a-nutshell/>

- Derivations
- Types of hybrid coordinates
- Nonlinear Coriolis term
- Kinetic energy gradient
- Explicit spatial filters
- Time discretization (split-explicit RK3); treatment of acoustic modes
- Top-of-model Rayleigh damping

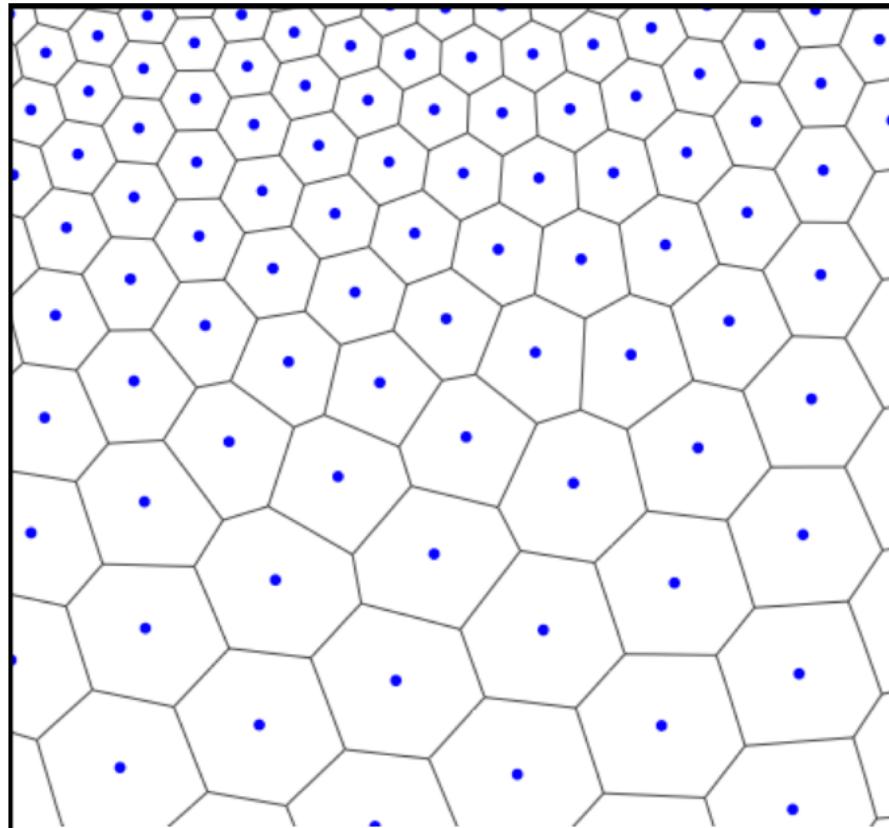
Summary

- MPAS is a cool model
- Hybrid-height, terrain-influenced vertical coordinate
- C-Grid staggering on a Voronoi mesh
- Nonhydrostatic, conservative equations
- Sophisticated FV discretization on an unstructured mesh



<http://mpas-dev.github.io/atmosphere/tutorial.html>

Thank you (???)



<http://mpas-dev.github.io/atmosphere/tutorial.html>

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