

A couple more things about strain:

(1) strain rate $\dot{\epsilon}$

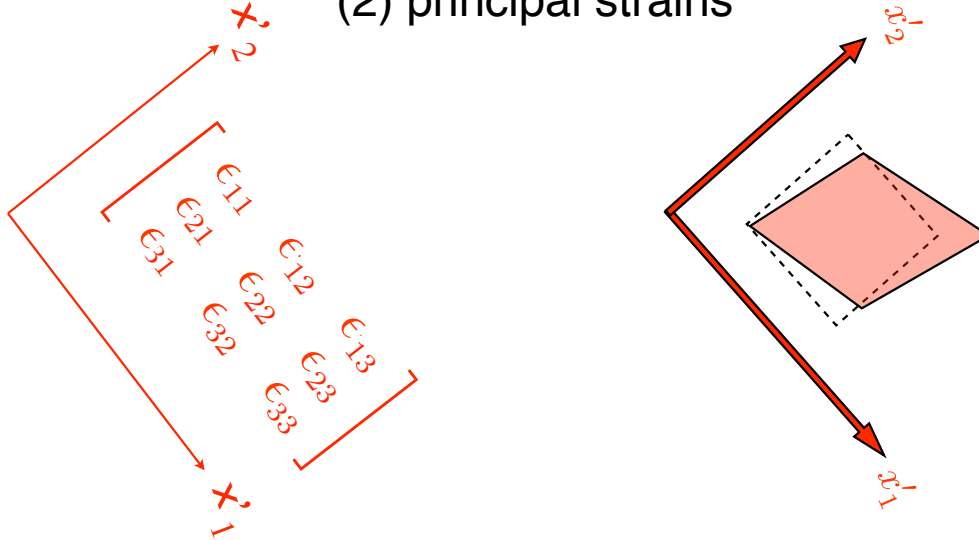
- strain per unit time
- denoted with a dot
- units?

$\dot{\epsilon}$ is obtained from derivatives of velocity (not displacement) with respect to position

$$\begin{aligned}
 & \begin{matrix} \left[\dot{\epsilon} \right] & \left[\dot{\omega} \right] \end{matrix} \\
 & \begin{pmatrix} \frac{\partial V_x}{\partial x} & \frac{\partial V_x}{\partial y} \\ \frac{\partial V_y}{\partial x} & \frac{\partial V_y}{\partial y} \end{pmatrix} = \begin{pmatrix} \frac{\partial V_x}{\partial x} & \frac{1}{2} \left(\frac{\partial V_x}{\partial y} + \frac{\partial V_y}{\partial x} \right) \\ \frac{1}{2} \left(\frac{\partial V_x}{\partial y} + \frac{\partial V_y}{\partial x} \right) & \frac{\partial V_y}{\partial y} \end{pmatrix} + \begin{pmatrix} 0 & \frac{1}{2} \left(\frac{\partial V_x}{\partial y} - \frac{\partial V_y}{\partial x} \right) \\ -\frac{1}{2} \left(\frac{\partial V_x}{\partial y} - \frac{\partial V_y}{\partial x} \right) & 0 \end{pmatrix} \\
 & \qquad \qquad \qquad \text{(symmetric)} \qquad \qquad \qquad \text{(anti-symmetric)} \\
 & \begin{pmatrix} \square \rightarrow & \square \nearrow \\ \square \nwarrow & \square \rightarrow \end{pmatrix} = \begin{pmatrix} \square \rightarrow & \square \nearrow \\ \square \nwarrow & \square \rightarrow \end{pmatrix} + \begin{pmatrix} 0 & \square \nearrow \\ \square \nwarrow & 0 \end{pmatrix} \\
 & \qquad \qquad \qquad = \begin{pmatrix} \dot{\epsilon}_{xx} & \dot{\epsilon}_{xy} \\ \dot{\epsilon}_{xy} & \dot{\epsilon}_{yy} \end{pmatrix} + \begin{pmatrix} 0 & \dot{\omega} \\ -\dot{\omega} & 0 \end{pmatrix} \quad \leftarrow \begin{array}{l} \text{Where the dot over a term} \\ \text{indicates a time derivative} \\ \text{and} \\ \omega \text{ is amount of rotation.} \end{array} \\
 & \qquad \qquad \qquad \text{strain rate} \qquad \qquad \qquad \text{rotation rate} \\
 & \qquad \qquad \qquad \text{matrix} \qquad \qquad \qquad \text{matrix} \\
 & \text{velocity} \Rightarrow \text{deformation} \quad \text{and} \quad \text{rotation} \\
 & \text{gradients} \quad \text{rate} \qquad \qquad \text{rate} \\
 & \qquad \qquad \qquad \left[\dot{\epsilon} \right] \qquad \qquad \qquad \left[\dot{\omega} \right]
 \end{aligned}$$

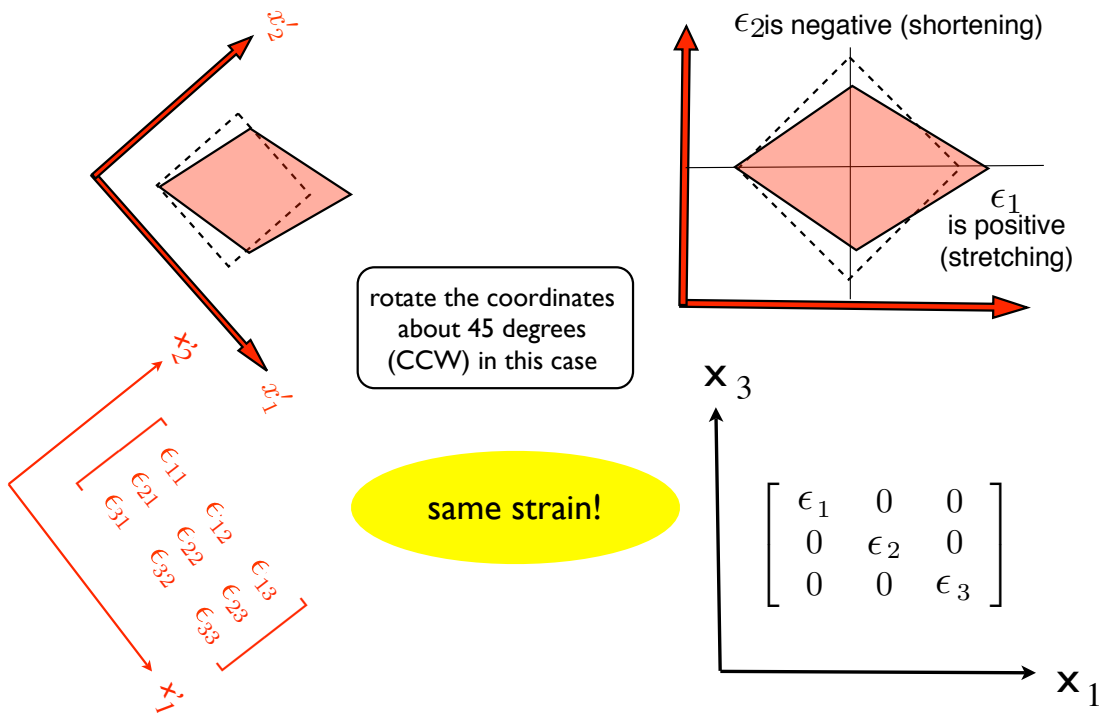
A couple more things about strain

(2) principal strains



suppose you know the strain matrix,
which was figured out in the (x'_1, x'_2, x'_3)
coordinate system

By rotating the coordinate system you can express 2D
strain with just two normal strains (“**principal strains**”)



Why does this matter? It becomes clear later on, when we try to apply Hooke's Law in the Earth, without resorting to linear algebra.

$$F = k \text{ times "x"} = (\text{constant}) \times (\text{spring displacement})$$



$$\text{"x"} = \epsilon L$$

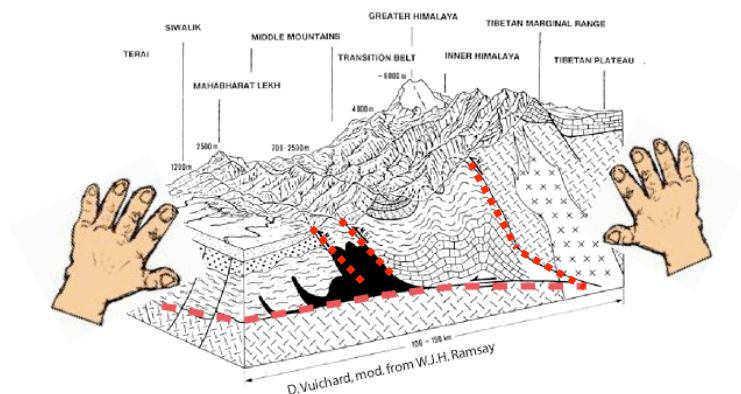
$$F = (kL)\epsilon = (\text{constant}) \times (\text{spring strain } \epsilon)$$

(where L is the length of the spring)

$$\text{Stress} = \sigma = (\text{constant}) \times (\text{strain})$$

Stress is just **force per unit area acting on a surface**

It is just "some constant" times strain



This means that stress is a tensor (that is, a matrix), too.

What are the units of force / area?

Normal stress causes normal strain

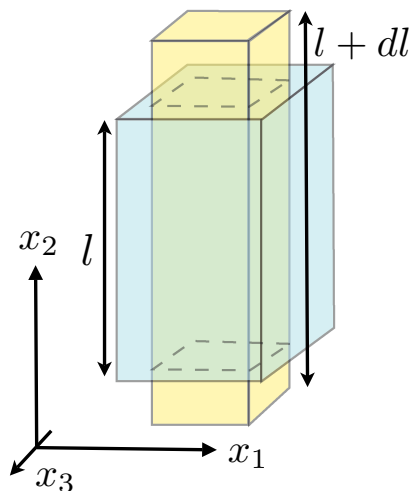
What are typical normal stresses in the Earth? How “big” is a Pascal?

Sign convention.

The mean of the three normal stresses is (-pressure).

Lithostatic pressure in the Earth...

Strain in the Earth due to normal stress:
Hooke's Law



$$\epsilon_{22} = \frac{dl}{l} = \frac{du_2}{dx_2}$$

$$\begin{aligned}\sigma_{22} &= E\epsilon_{22} \\ \epsilon_{22} &= \frac{\sigma_{22}}{E}\end{aligned}$$

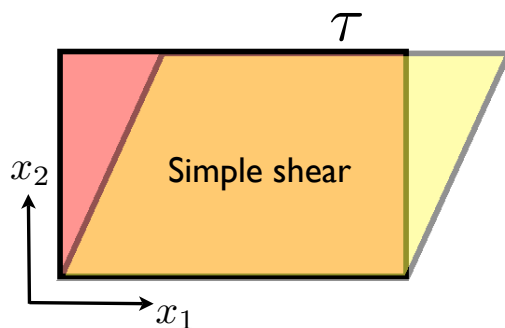
What is E?

E is the “Young’s Modulus”

In rocks, E is very, very large.

typical value is 7.5×10^{10} Pascals

Shear stress (sometimes called τ)
causes shear strain



$$\begin{aligned}\tau_{21} &= 2G\epsilon_{21} \\ \epsilon_{21} &= \frac{\tau_{21}}{2G}\end{aligned}$$

G is the “Shear Modulus”

In rocks, G is also very, very large.

typical value is 3×10^{10} Pascals