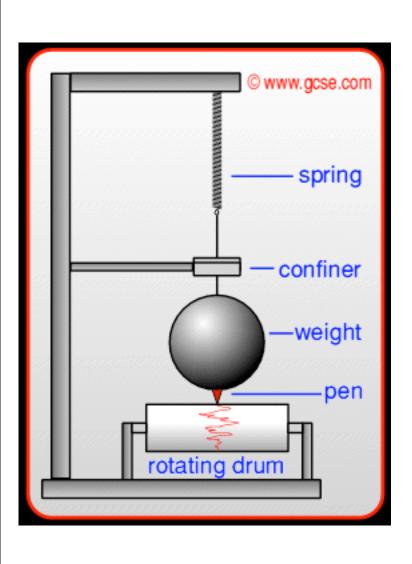
SEISMOMETERS, SEISMOGRAPHS, SEISMOGRAMS

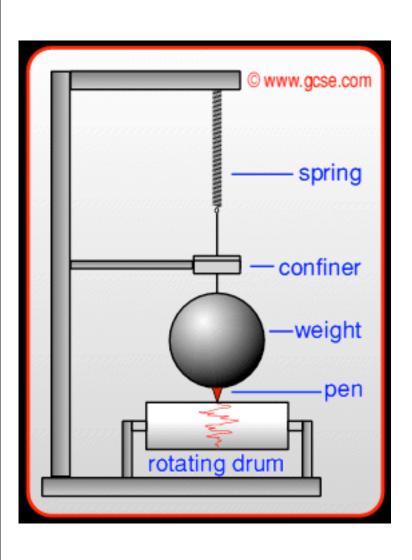
- I. What is a seismometer?
- 2. What is a seismograph?
- 3. What is a seismogram

A **seismometer** is a mechanical device that measures and amplifies ground motion at a point on the Earth's surface or in a borehole



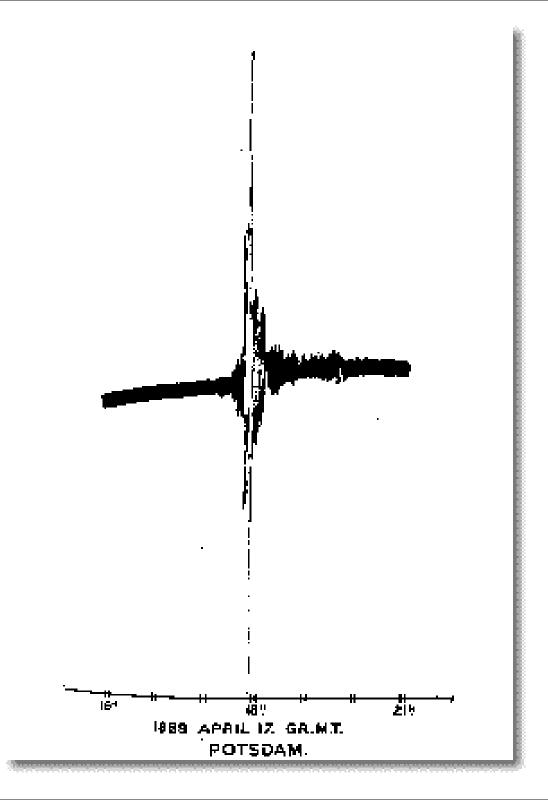


A modern **seismograph** records ground motion (from a seismometer) in digital format onto magnetic or optical disk





A **seismogram** is a visual representation of ground motion at a point in space as a function of time



SEISMOMETERS MEASURE GROUND MOTIONS

- > ground motions can be described and measured in different ways:
- I. ground displacement
- 2. ground velocity
- 3. ground acceleration

- Q1. How are they related?
- Q2. Which is most useful?

displacement

u(t)

velocity

 $\frac{du(t)}{dt}$

acceleration

 $\frac{d^2u(t)}{dt^2}$

damage ~ force ~ acceleration

During large earthquakes, accelerations can approach or even exceed gravity

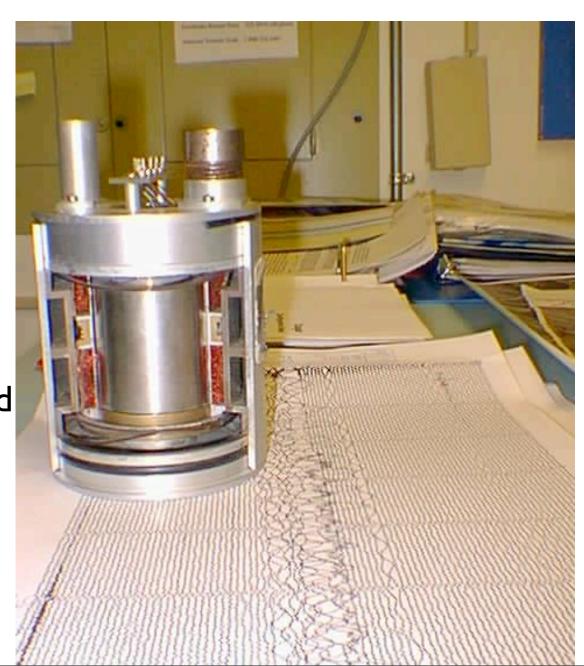
SEISMOMETRY EXERCISE

- > ground motions provide much important information on both earthquakes and Earth structure
- >NO seismometer provides a perfect representation of ground motion, each one has an (imperfect) **response**
- > we will derive response for a simple damped pendulum seismometer
- > GROUP EXERCISE: I want you to analyse this response to see how true ground motions are modified by seismometer

SHORT/LONG PERIOD SEISMOMETERS & GEOPHONES

- > used prior to 1990's
- > work on damped pendulum theory
- resonant frequency at I Hz,0.1 Hz
- > mass incorporates solenoid which moves in a magnetic field
- > Faraday's law states

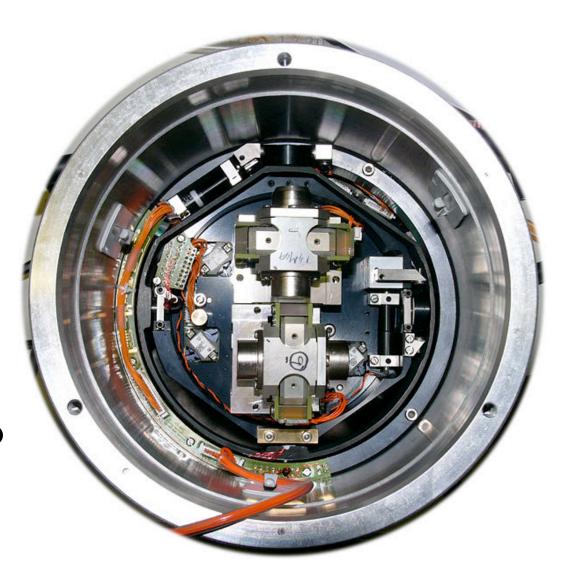
$$\epsilon = -\frac{d\Phi}{dt} \sim \frac{dv}{dt}$$



MODERN BROADBAND SEISMOMETERS

- > record motions faithfully between 100 0.001 Hz
- > driven by sophisticated feedback electronic circuits

> motion is measured through voltage required to keep masses stationary



STRONG MOTION SEISMOGRAPHS

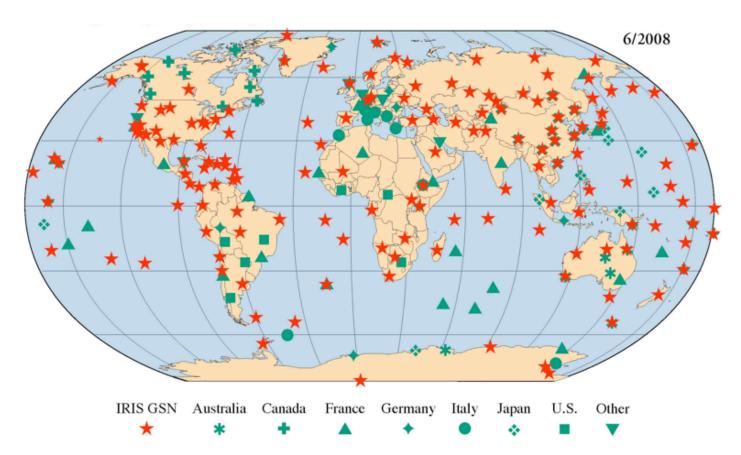
- > made from MEMS & sensitive to large accelerations
- > regular seismometers go off scale
- > used in triggered mode to study effects of large eq's
- > employed by engineers to aid in design of earthquake resistant infrastructure



SEISMIC NETWORKS

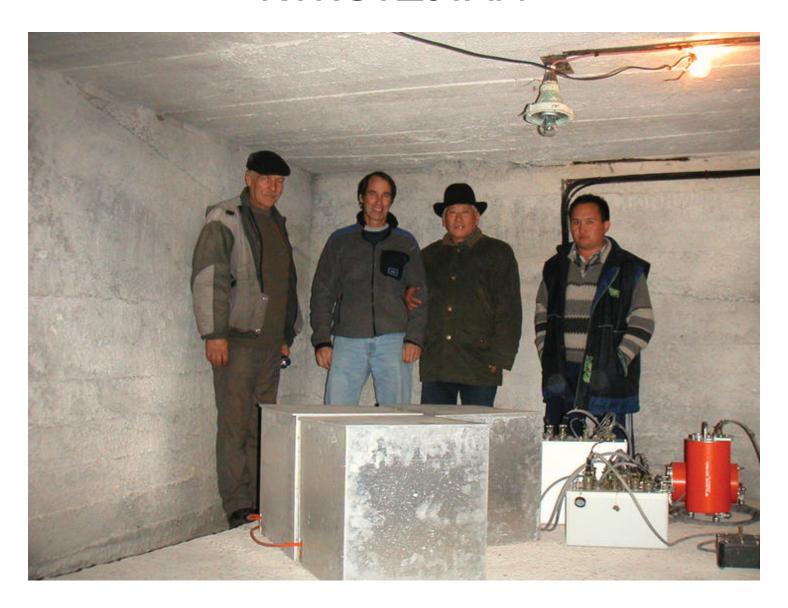
- > arrays of seismometers deployed for a common purpose
- I. Global Seismic Network
- 2. Regional Networks
- 3. Portable Arrays
- 4. EarthScope

GLOBAL SEISMIC NETWORKS



- > 150+ stations globally distributed
- > high quality stations with detection limit ~M=4
- > partly underwritten by military agencies to aid in nuclear test ban verification treaties

UNDER GROUND VAULT - KYRGYZSTAN



> note thermal insulation, concrete bunker

SOUTH POLE SITE



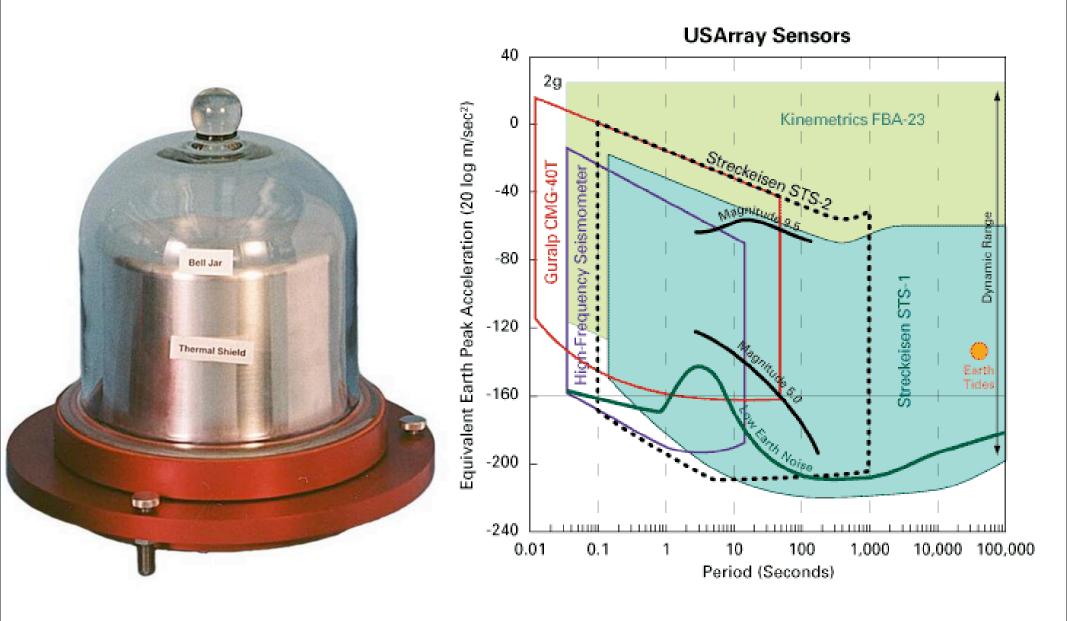
> some sites involve seismometers in boreholes to minimize noise

PITCAIRN ISLAND SITE



> most communications by satellite

HIGH PRECISION SEISMOMETERS



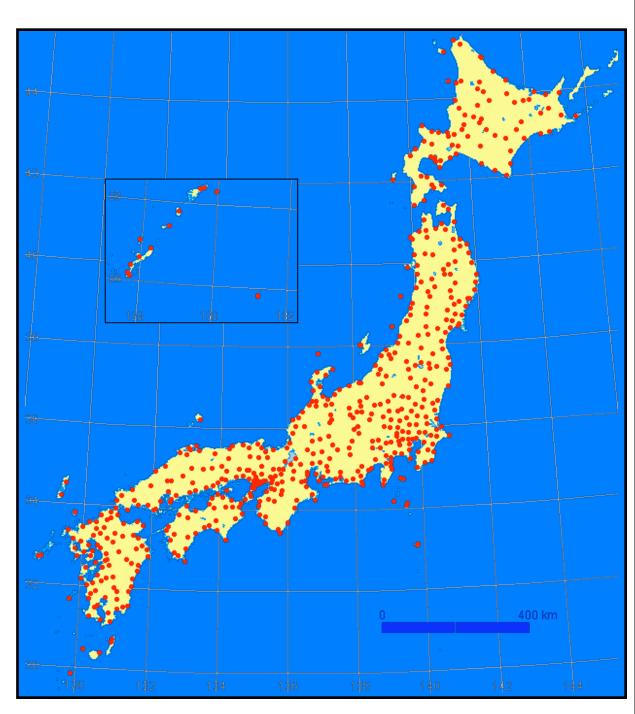
> highly sensitive : -200 is equivalent to what acceleration?

$$x = 10^{-10} \text{m s}^{-2}$$

REGIONAL SEISMOGRAPH NETWORKS

- > Japanese Hi-Net has over 600 short-period, borehole stations
- > since 2000, has led to many important discoveries
- > 10-20 km spacing

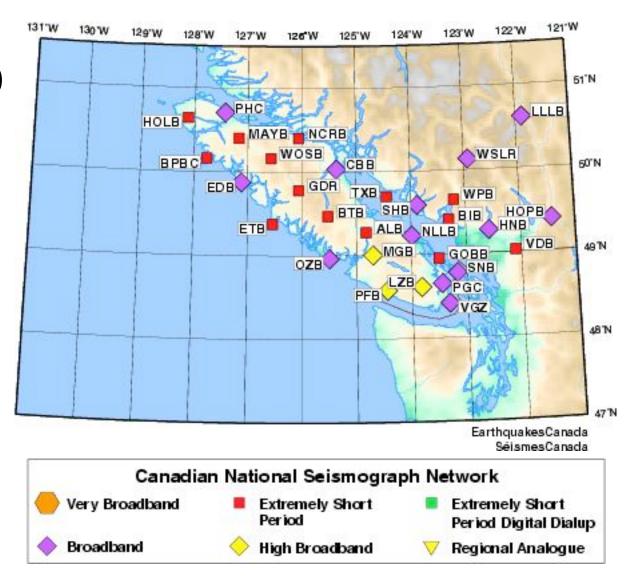




CANADIAN NATIONAL SEISMOGRAPH NETWORK (B.C.)

> G.S.C. operates ~30 seismographs in SW B.C.

> note concentration on V.I. and lower mainland



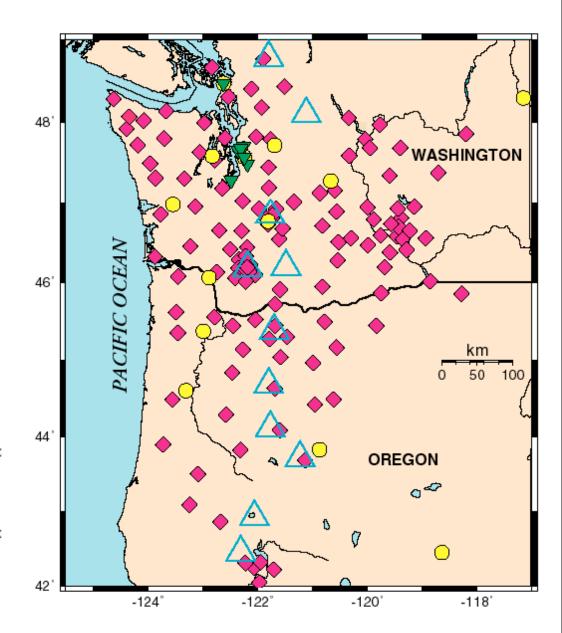
PACIFIC NORTHWEST SEISMIC NETWORK

> UW operates ~100 sp and ~10-20 BB sites through Washington and Oregon

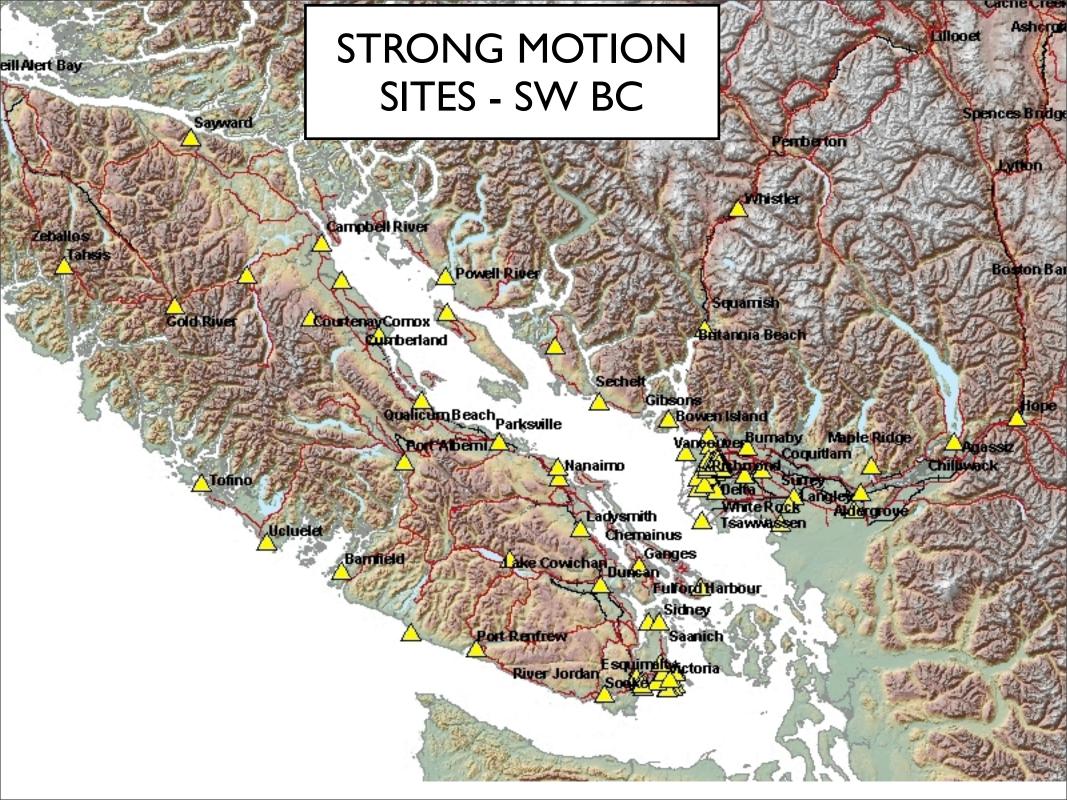
> significant data exchange between CNSN and PNSN Accelerometer: Measures

Accelerometer: Measures strong ground motion.

- Seismometer (3 component):
 Measures vertical and N-S
 and E-W ground motions.
- Seismometer (1 component): Measures only vertical ground motions.





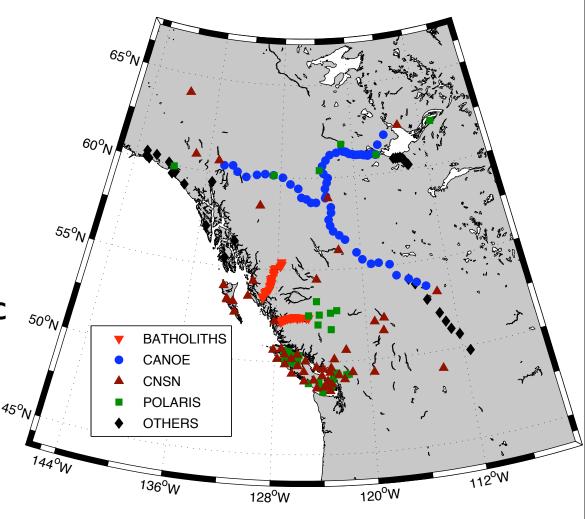


PORTABLE ARRAYS

> many countries possess portable instruments used for temporary field campaigns

Canada: POLARIS (Portable
 Observatories for Lithospheric
 Analysis and Research
 Investigating Seismicity

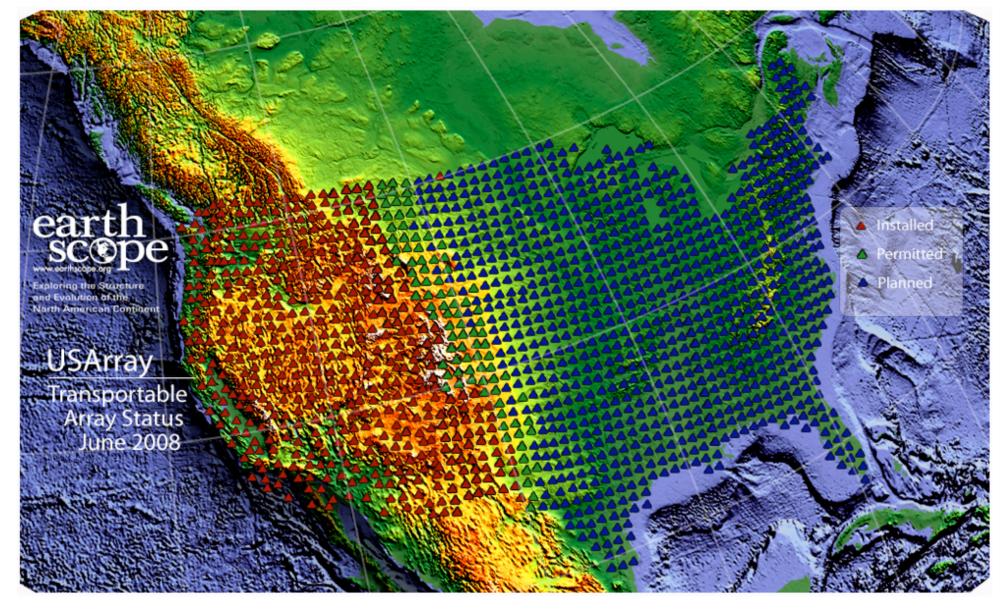
> can be used in aftershock or structural studies



PORTABLE ARRAY VAULTS

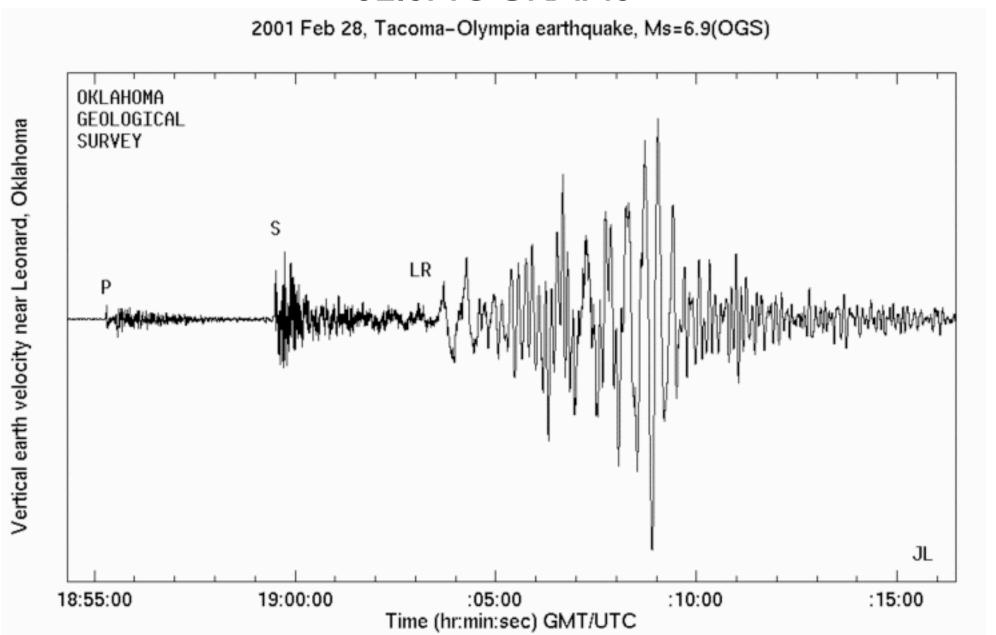
- > makeshift vaults with solar power
- > data archived onto loggers that record continuously
- typical deploymentI-2 years



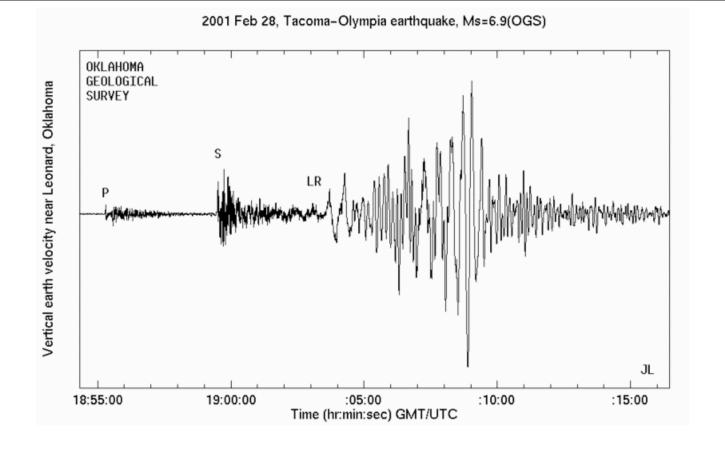


- > new generation of portable experiment; cover whole USA at 70 km spacing
- > each station active for 18 months, deployed roll-along array over 15 years

SEISMOGRAMS



> incredibly rich and varied in appearance depending on source, frequency content, distance etc.



Many ways to extract information from seismogram:

- I. Identify main phases, extract time/amplitude
- 2. Identify scattered phases
- 3. Match whole seismogram

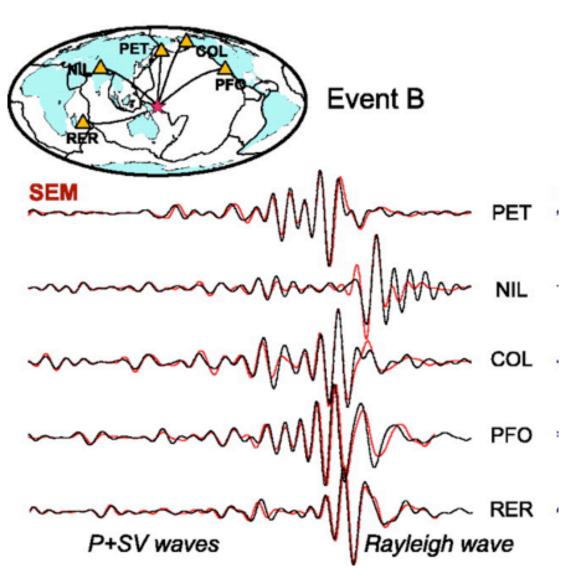
Use information to extract knowledge of **earthquake** and/or **earth structure**

LONG PERIOD SEISMOGRAMS

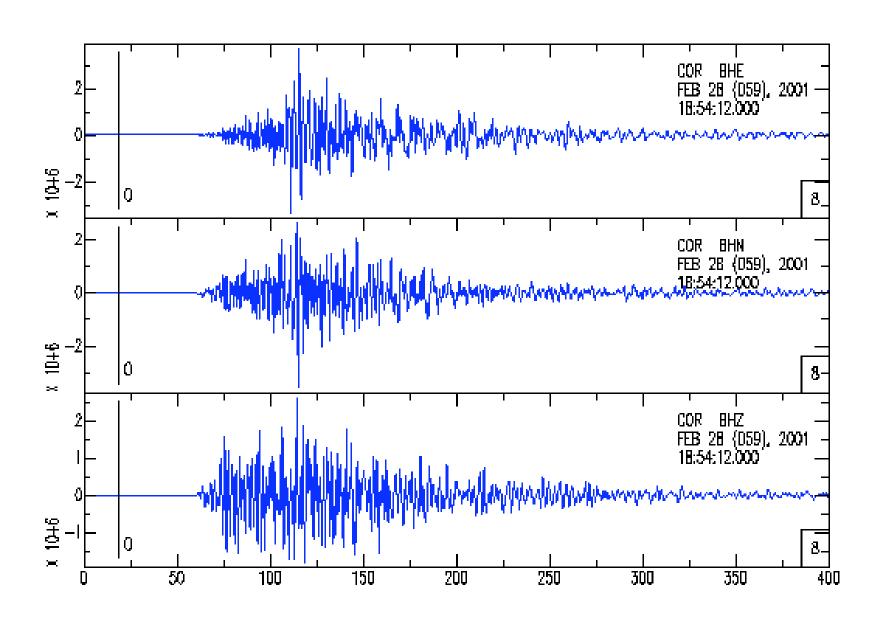
> T > 10 s

> dominated byS and surface waves

> simple to model

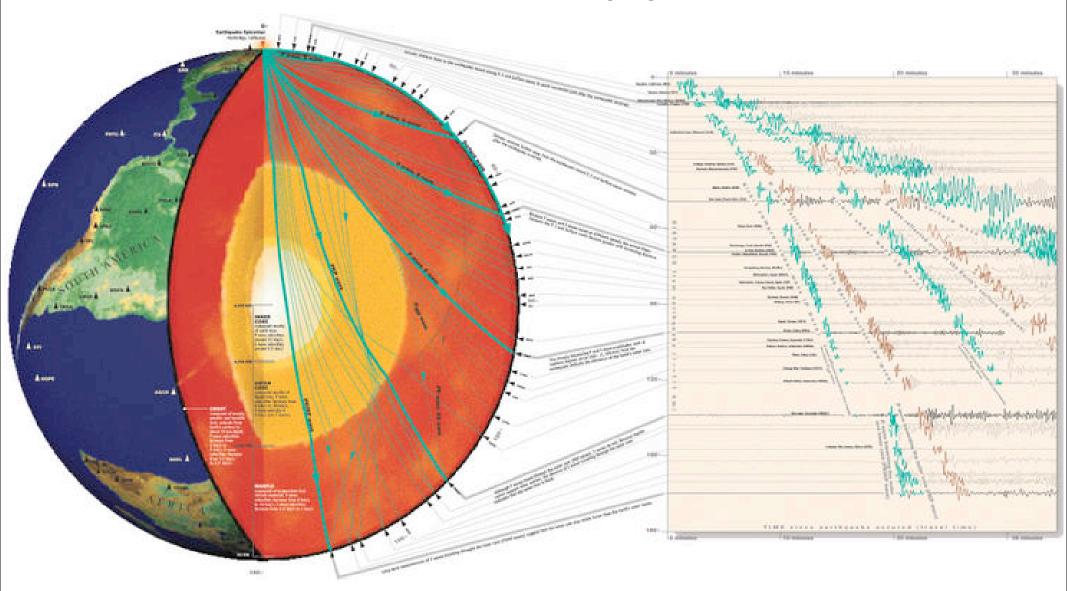


REGIONAL SEISMOGRAMS



> higher frequency (> I Hz), more complex, harder to identify individual P, S, surface waves

PRIMARY PHASES



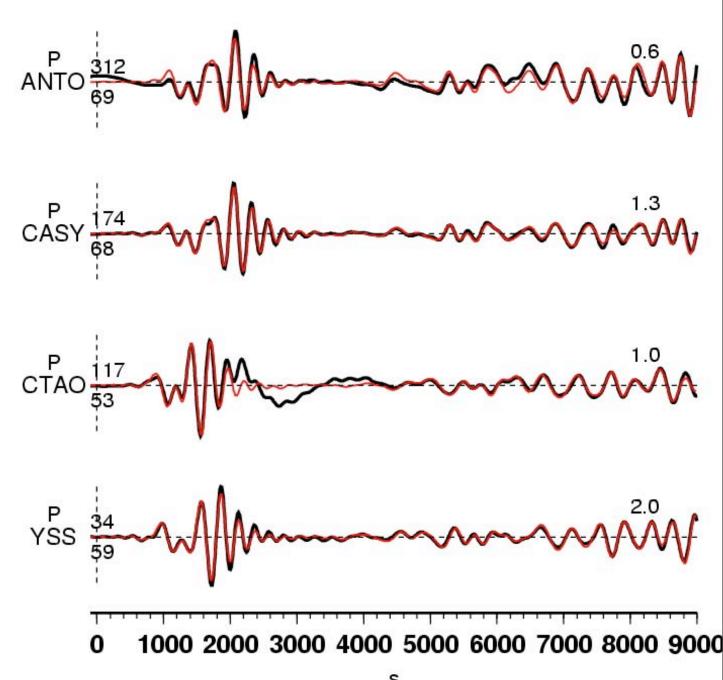
- > at global scale Earth looks like layered sphere
- > readily predictable seismic phases propagating through crust, mantle, outer core, inner core

WAVEFORM MATCHING

> requires knowledge knowledge of model for both source and struture

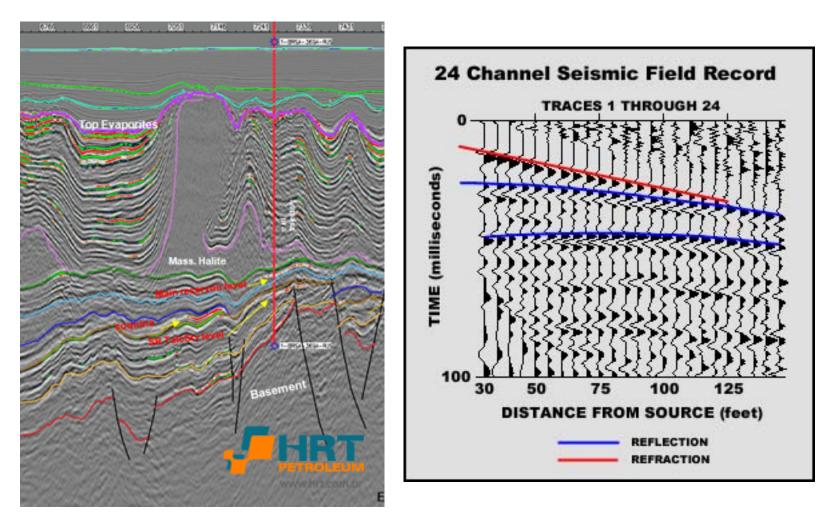
> data: black, model red

> at long periods (gt 10 s) we can model seismograms well, not so at short periods (lt 1s)



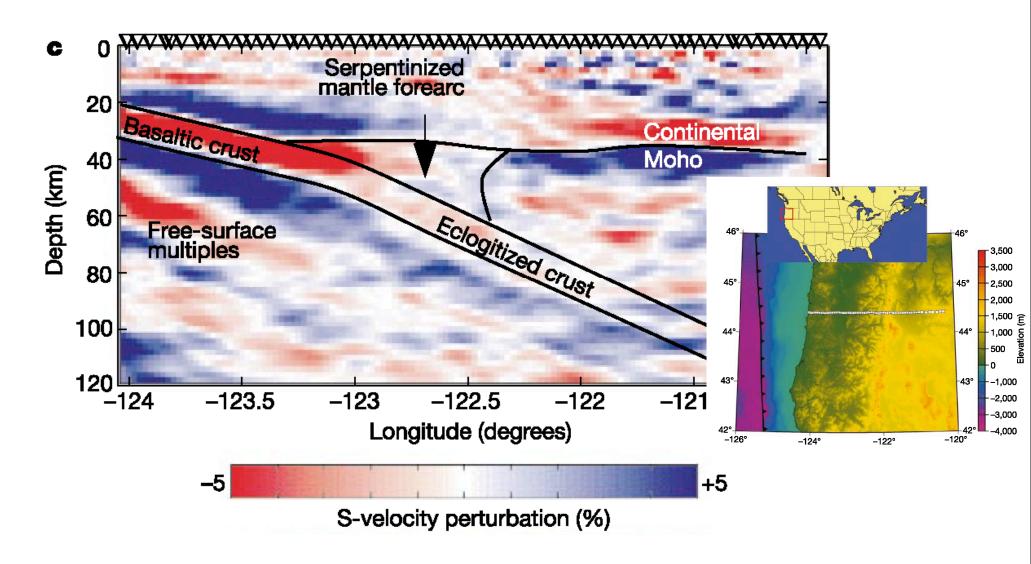
05/03/28 (Mw 8.5), Sumatra Earthquake

SEISMIC REFLECTION IMAGING



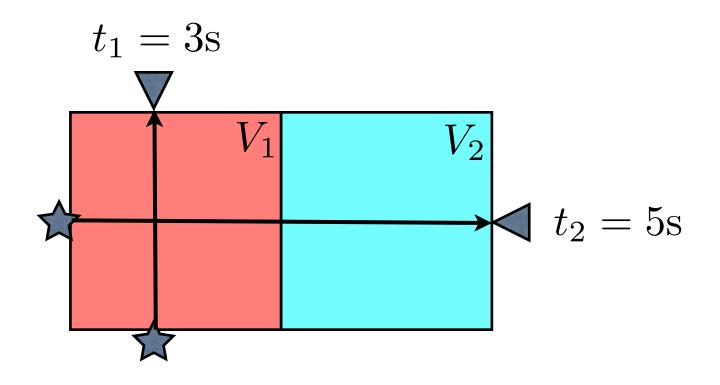
- > seismic reflections/conversions originate from discontinuities in velocity/density
- > primary tool for hydrocarbon exploration

TELESEISMIC IMAGING



> waves generated by earthquakes and scattered from discontinuties can be used to image e.g. subduction zones

TRAVELTIME TOMOGRAPHY



- > 2 blocks (I m X I m) with unknown velocities V_1, V_2
- > 2 traveltime measurements t_1, t_2
- > can you determine V_1, V_2 ?

SOLUTION

- I. Solve for V_1 using t_1 : $V_1 = 0.3333 \, \text{m/s}$
- 2. Insert in solve for V_2 : $V_2 = 0.5 \,\mathrm{m/s}$

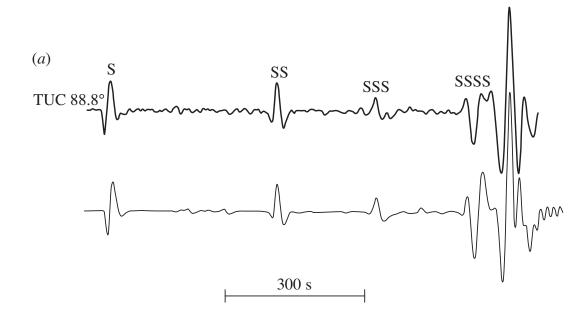
OR

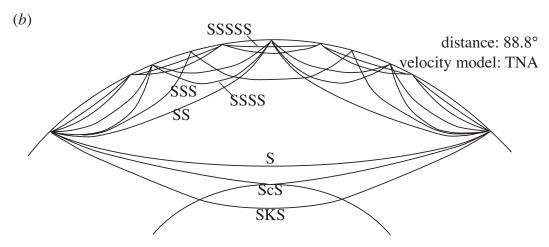
Solve matrix equation

$$\left[\begin{array}{c} t_1 \\ t_2 \end{array}\right] = \left[\begin{array}{c} 3 \\ 5 \end{array}\right] = \left[\begin{array}{c} 1 & 0 \\ 1 & 1 \end{array}\right] \left[\begin{array}{c} \frac{1}{V_1} \\ \frac{1}{V_2} \end{array}\right]$$

TRAVELTIMES OF MAJOR PHASES

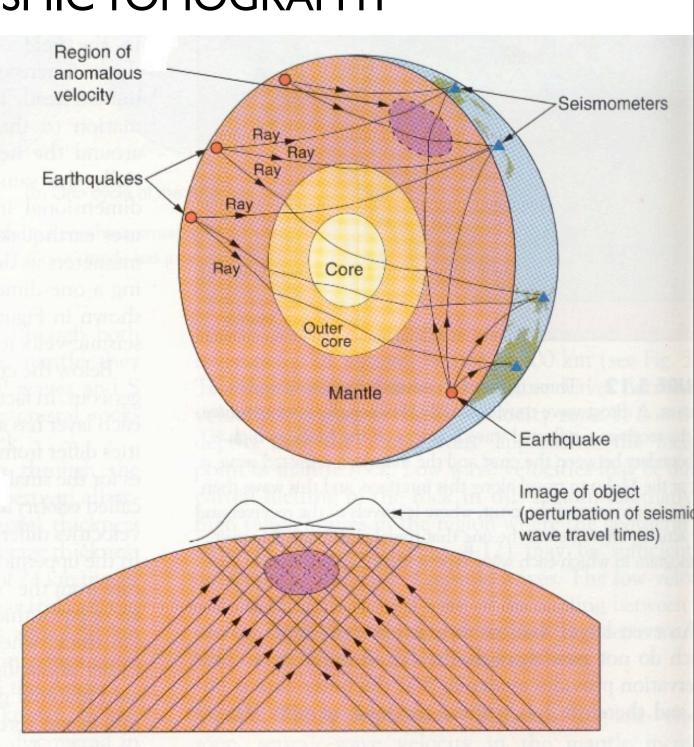
> modern global tomography incorporates information from many phases including CMB/ surface reflections

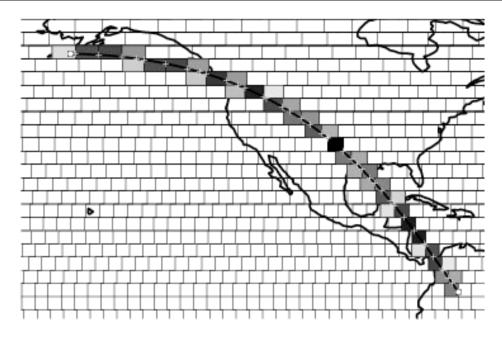




GLOBAL SEISMIC TOMOGRAPHY

- > model Earth as many constant velocity elements
- > each measured time represents a ray and an equation
- > 1,000,000's of equations in 100,000's of elements
- > solve enormous matrix system for Earth's velocity structure





Discretization: model earth as a mesh of small elements each with constant seismic velocity

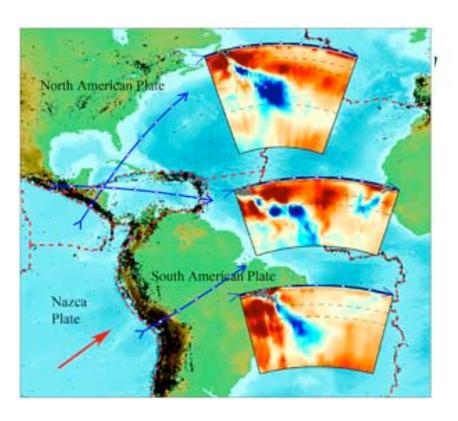
$$\delta t_i = L_{ij} \delta s_j$$
 or $\delta \mathbf{t} = \mathbf{L} \cdot \delta \mathbf{s}$

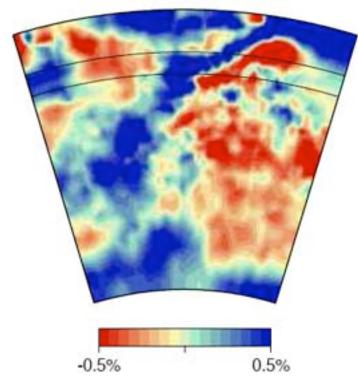
M travel-time anomalies
$$\begin{vmatrix} \vdots \\ \delta t_i \\ \vdots \end{vmatrix} = \begin{bmatrix} \vdots \\ \dots \\ L_{ij} \\ \vdots \end{vmatrix} \times \begin{bmatrix} \vdots \\ \delta s_j \\ \vdots \end{bmatrix}$$
N slowness perturbations

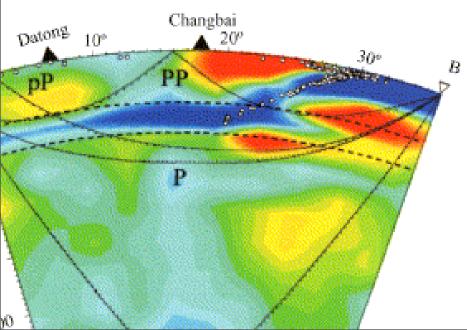
M×N sensitivity matrix

WHAT DOES TOMOGRAPHY TELL US ABOUT PLATE BOUNDARIES AND MAJOR FAULTS?

TOMOGRAPHIC IMAGES







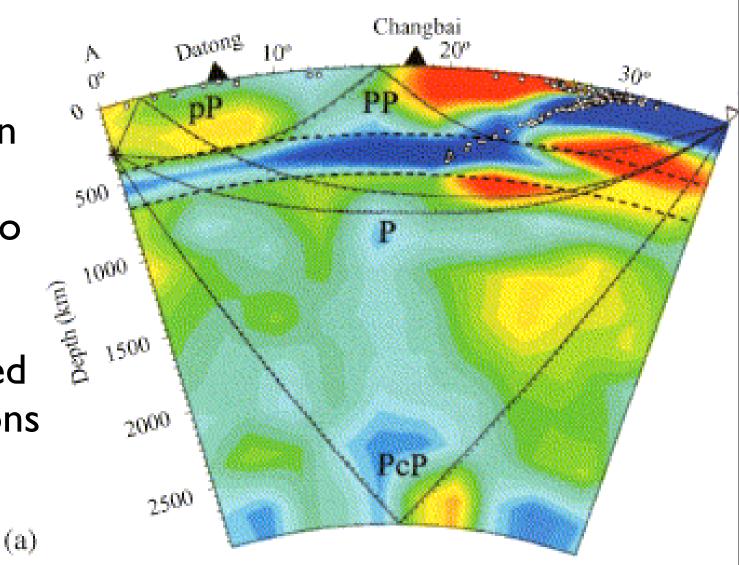
on zones as high-velocity bugh the mantle to the CMB

ocity?

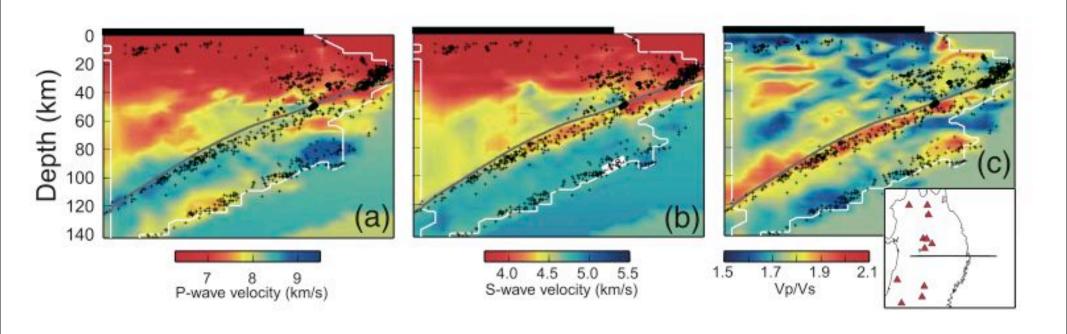
TOMOGRAPHY & WADATI-BENIOFF ZONES

> most subduction zones have W-B zones extending to <700 km depth</p>

> seismicity related to mineral reactions



DOUBLE BENIOFF ZONES



- > example from NE Japan
- > 2 lines of seismicity, one near plate boundary, the other 40 km below
- > thought to be due to dehydration reactions

CASCADIA SEISMICITY

Velocity (m s⁻¹)

- > Cascadia not as active as some s.z.'s
- > still controversy over exact position of downgoing plate
- > W-B seismicity restricted to above 100 km

50-

20 km

